A Method for Estimating Volume and Rate of Runoff in Small Watersheds

ABSTRACT

The Soil Conservation Service (SCS) has developed charts ES-1026 and ES-1027 for estimating the instantaneous peak discharge expected from small areas. They provide the peak discharge rate for establishing conservation practices on individual farms and ranches and for the design of water-control measures in small watersheds. The graphs were prepared from computations made by automatic data processing (ADP). Each graph relates peak discharge to drainage area and rainfall depths for each of (1) a given set of watershed characteristics, (2) different rainfall time distributions and (3) three categories of average watershed slopes. Peak discharges range from 5 to 2,000 cubic feet per second (cfs), drainage areas range from 5 to 2,000 acres, and 24-hour rainfall depths range from 1 to 12 inches. Curve numbers (CN) are used to represent watershed characteristics that influence runoff. Each chart represents one of seven curve numbers ranging from 60 to 90 in increments of 5. Each group of seven charts represents one of the three average watershed slope factors (FLAT, MODER-ATE, and STEEP) making a total of 21 charts for each of two rainfall time distributions. The procedures for computation of peak discharges by ADP were based upon those in the SCS National Engineering Handbook, Section 4, Hydrology, August 1972. The logic and procedures used for the ADP computation are described.

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A Method for Estimating Volume and Rate of Runoff in Small Watersheds

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INTRODUCTION

Ven Te Chow has described many methods which have been used for determining waterway areas and the design of drainage control structures in small watersheds ($\underline{1}$). Some of these methods have been used by the Soil Conservation Service (SCS) for estimating peak discharge rates. These include the rational method (Ramser curves after C. E. Ramser), the Cook method after H. L. Cook, the modified Cook or EW method by M. M. Culp and others, and the method by Victor Mockus and others described in the National Engineering Handbook, Section 4, Hydrology (NEH-4) an inservice handbook of SCS ($\underline{7}$). SCS has used these methods primarily for the design of measures for individual farms and ranches.

The NEH-4 method provides for the development of a complete hydrograph and involves more detailed computations than the others. It is used primarily for planning and designing larger measures—larger than those for farms and ranches—in watersheds planned under the Water—shed Protection and Flood Prevention Act (Public Law 566, 83d Cong.; Stat. 666), as amended.

Using different methods under similar conditions SCS, obtained wide differences in the peak rates. These differences were mainly due to the choice of coefficients and factors inherent in each method rather than to the method itself. The method adopted by SCS is shown in charts ES-1026 and ES-1027 (appendix). Guidelines have been established for selecting nationally applicable values for this method's parameters. This set of parameters is expected to provide adequate and more uniform estimates of peak discharges between areas having similar watershed characteristics.

A primary requirement was that the method be simple enough to be used by all grades of professional and subprofessional personnel in SCS. They all need to make quick, on-the-spot estimates of peak discharge rates for planning and designing soil and water conservation measures.

It is further desirable for the method to be closely allied with those in NEH-4. The peak discharge for a small watershed with unusual characteristics can then be computed using the more detailed procedures in NEH-4 but with the same parameters. Specific values are computed for each parameter in contrast to the average values used in the charts.

The method described here is generally limited to drainage areas of 2,000 acres or less and to watersheds that have average slopes of less than 30 percent. The NEH-4 method is generally used for watersheds exceeding these limits or when the computed peak discharge exceeds 2,000 cfs. There are other circumstances where the method described here may not provide adequate estimates and the NEH-4 method should be used. These are described later under pertinent headings.

STORM RAINFALL

Stream-gage measurements are rarely available for small watersheds. Generalized rainfall data, however, are available nationally. Therefore it is desirable that the national SCS method for computing peak discharge rates and runoff volumes in small areas use rainfall for their basic input.

The Weather Bureau's Rainfall-Frequency Atlases covering the United States, Puerto Rico, and the Virgin Islands provide rainfall-frequency data for areas less than 400 square miles, for durations to 24 hours, and for frequencies from 1 to 100 years (6, 8, 9, 10, 11).

Adjustment of rainfall with respect to area is

Adjustment of rainfall with respect to area is not necessary in the method described because the drainage areas are small. But the distribution of storm rainfall with respect to time is an important parameter. Two major regions were identified for this purpose. Time distributions for each are tabulated in table 1 and shown in figure 1. Type I represents regions with a maritime climate. Type II represents regions in which the high rates of runoff from small areas are usually generated from summer thunderstorms.

The type I and type II distributions are based on generalized rainfall depth-duration relationships obtained from Weather Bureau technical papers. The accumulative graphs in figure 2, which are the basis for type I and II distributions, were established by (1) plotting a ratio of rainfall amount for any duration to the 24hour amount against duration for a number of locations and (2) selecting a curve of best fit. Selected curves are shown as dashed lines in figure 2. Note that the type II distribution (fig. 2) underestimates the 1-hour duration by about 0.6 inch at Lincoln, Nebr., overestimates it by about 0.5 inch at Mobile, Ala., and is within O.l inch on the northwest corner of Utah. The type I distribution underestimates the 6-hour duration by about 1 inch at Kahuka Point, Oahu, Hawaii. These variations are within the accuracy of rainfall amounts read from the Weather Bureau references.



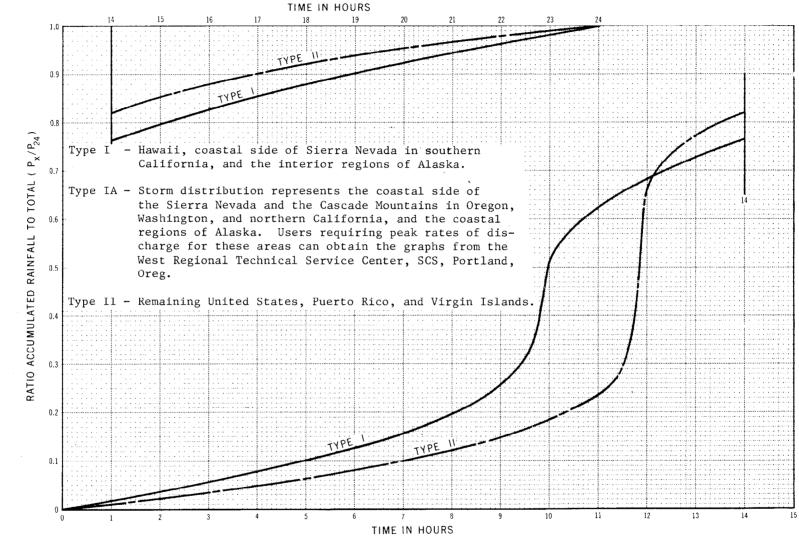


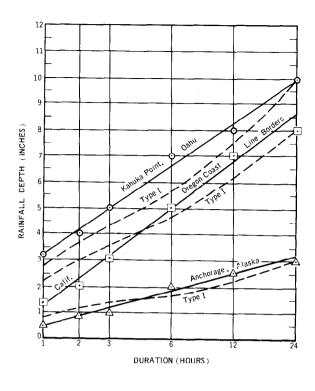
Figure 1.--Twenty-four-hour rainfall distributions (SCS).

Time	P _X /F	241/
(hours)	Type I	Type II
0 2.0 4.0 6.0 7.0 8.0 8.5 9.0 9.5 9.75 10.0 10.5 11.75 12.0 12.5 13.0 13.5 14.0 16.0 20.0 24.0	0 .035 .076 .125 .156 .194 .219 .254 .303 .362 .515 .583 .624 .654 .682 .727 .767 .830 .926 1.000	0 .022 .048 .080 .120 -147 .163 .181 .204 .235 .283 .387 .663 .735 .772 .799 .820 .880 .952 1.000

1/Ratio accumulated rainfall to total.

Average intensity-duration values used to develop the dashed lines in figure 2 are rearranged to form the type I and II distributions in figure 1. The type I distribution is arranged so that the greatest 30-minute depth occurs at about the 10-hour point of the 24-hour period, the second largest in the next 30 minutes, and the third largest in the preceeding 30 minutes. This alternation continues with each decreasing order of magnitude until the smallest increments fall at the beginning and end of the 24-hour rainfall (fig. 1). The type II distribution is arranged in a similar manner but the greatest 30-minute depth occurs near the middle of the 24-hour period. The selection of the period of maximum intensity for both distributions was based on design consideration rather than meteorological factors.

The effective storm period that contributes to an instantaneous peak rate of discharge varies with the time of concentration $(T_{\rm c})$ of each small watershed. It is only a few minutes for a very short $T_{\rm c}$ and up to 24 hours for a long $T_{\rm c}.$ The effective period for most watersheds smaller than 2,000 acres is less than 6 hours. Because of the "built-in" range of 30-minute intensities the 24-hour duration is equally appropriate for a 5-acre watershed with less than a 30-minute



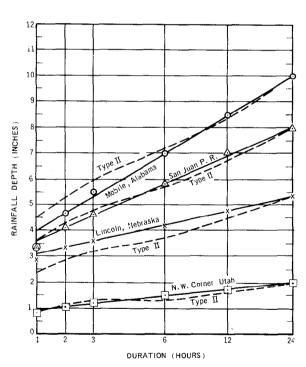


Figure 2.—Generalized 25-year frequency rainfall depth-duration relationships (U.S. Weather Bureau Rainfall Atlases).

effective storm period as it is for a 2,000-acre watershed where the effective periods may take up the entire 24 hours.

RAINFALL-RUNOFF EQUATION

The runoff equation used by SCS was developed by Victor Mockus and others about 1947 (2, 5, 7). A relationship between accumulated rainfall and accumulated runoff was derived from experimental plots for numerous soils and vegetative cover conditions. Data for land-treatment measures, such as contouring and terracing, from experimental watersheds were included. The equation was developed mainly for small watersheds for which only daily rainfall and watershed data are ordinarily available. It was developed from recorded storm data that included total amount of rainfall in a calendar day but not its distribution with respect to time. The SCS runoff equation is therefore a method of estimating direct runoff from storm rainfall of 1 day or less.

The equation is:

$$Q = \frac{(P - I_a)^2}{(P - I_a) + S} \tag{1}$$

Where:

Q = accumulated direct runoff.

P = accumulated rainfall (potential maximum runoff).

 ${\rm I_a}$ = initial abstraction including surface storage, interception, and infiltration prior to runoff.

S = potential maximum retention.

The inset in figure 3 shows the initial abstraction (I_a) in a typical storm. The relationship between I_a and S was developed from experimental watershed data. It removes the necessity for estimating I_a for common usage. The empirical relationship used in the SCS runoff equation is:

$$I_8 = 0.2S$$
 (2)

Substituting 0.2S for I_{a} in equation (1), the equation follows:

$$Q = \frac{(P - 0.2S)^2}{P + 0.8S}$$
 (3)

To show the rainfall-runoff relationship graphically, S values are transformed into curve

numbers (CN) by the following equation (fig. 3):

$$CN = \frac{1000}{10 + S} \tag{4}$$

The S values for CN's ranging from 0 to 100 are tabulated in NEH-4, table 10.1. Research data provided the association of CN's with various hydrologic soil-cover complexes as shown in table 2 for an average antecedent moisture condition. Soils are divided into four hydrologic soil groups: A, B, C, and D. Group A soils have a high infiltration rate even when thoroughly wet. When thoroughly wet, group B soils have a moderate infiltration rate, and group C soils a slow infiltration rate, and group D soils a very slow infiltration rate. Table 7.1 of NEH-4 lists more than 9,000 soils and their hydrologic group.

The rainfall-runoff chart (fig. 3) is used mostly for estimating the runoff from watersheds for which composite CN's are obtained from listings like those in table 2. The curves can in turn be used to estimate a composite CN for an unlisted watershed characteristic with rainfall and runoff data for only a few years. rainfall-runoff values for each storm in the short period can be plotted on a facsimile of figure 3. The curve in figure 3 equally dividing the plotted points can be assumed to represent the runoff CN for an average antecedent moisture condition in the watershed. The plotted points are usually widely scattered, representing a change in the value of S in equation (3) and hence a corresponding change in CN from one storm to the next. Most of this difference is the result of variations in soil moisture preceding each storm. Mockus based the antecedent moisture condition (AMC) on the total rainfall in the 5-day period preceding a storm and divided the AMC into three conditions (table

Figure 4 demonstrates how the plotted points usually fall between the CN's representing AMC I and AMC III with AMC II equally dividing them. This capability is an advantage to engineers working in foreign countries where, without experimental data on watershed characteristics unique to the local area, a minimum amount of measured data may suffice to establish CN's adequate for the design of small structures.

Changes in plant cover between seasons along with changes in land use from year to year can also affect the degree of scatter of plotted P and Q points. Furthermore, if rain gages are not spaced close enough to measure watershed precipitation accurately, this will cause unrealistic scatter in the P and Q plotting.

The peak discharge computations in ES-1026 and ES-1027 are based on AMC-II.

Figure 3.—Solution of the runoff equation, $Q = \frac{(P - 0.2S)^2}{P + 0.8S}$

Table 2.—Runoff curve numbers for hydrologic soil-cover complexes (Antecedent moisture condition II, and $\rm I_a$ = 0.2 S)

Land use and treatment	Hydrologic	Hyd	rologi	c soil	group
or practice	condition	A	В	С	D
Fallow					
Straight row		77	86	91	94
Row crops					
Straight row	Poor	72	81	88	91
Straight row	Good	67	78	85	89
Contoured	Poor	70	79	84	88
Contoured	Good	65	75	82	86
Contoured and terraced	Poor	66	74	80	82
Contoured and terraced	Good	62	71	78	81
Small grain					
Straight row	Poor	65	76	84	88
Straight row	Good	63	75	83	87
Contoured	Poor	63	74	82	85
Contoured	Good	61	73	81	84
Contoured and terraced	Poor	61	72	79	82
Contoured and terraced	Good	59	70	78	81
Close-seeded legumes or					
rotation meadow					
Straight row	${ t Poor}$	66	77	85	89
Straight row	Good	58	72	81	85
Contoured	${\tt Poor}$	64	75	83	85
Contoured	Good	55	69	78	83
Contoured and terraced	${ t Poor}$	63	73	80	83
Contoured and terraced	Good	51	67	76	80
Pasture or range					
No mechanical treatment	${\tt Poor}$	68	79	86	89
No mechanical treatment	${ t Fair}$	49	69	79	84
No mechanical treatment	Good	39	61	74	80
Contoured	${ t Poor}$	47	67	81	88
Contoured	Fair	25	59	75	83
Contoured	Good	6	35	70	79
Meadow	Good	30	58	71	78
Woods	Poor	45	66	77	83
	Fair	36	60	73	79
	Good	25	55	70	77
FarmsteadsRoads 1		59	74	82	86
Dirt		72	82	87	89
Hard surface	***	71	84	90	92

^{1/} Including rights-of-way.

Table 3.--Curve numbers (CN) for wet (AMC III) and dry (AMC I) antecedent moisture conditions corresponding to an average antecedent moisture condition (AMC II).

	ding CN's
AMC I	AMC III
100 87 78 70 63 57 51 45 40 35 31 26 22 18 15	100 98 96 94 91 88 85 82 78 74 70 65 60 55
12 9 6 4 2	43 37 30 22 13
	AMC I 100 87 78 70 63 57 51 45 40 35 31 26 22 18 15 12 9 6 4

1/ AMC I. Lowest runoff potential. Soils in the watershed are dry enough for satisfactory plowing or cultivation.

AMC II. The average condition.

AMC III. Highest runoff potential.

Soils in the watershed are practically saturated from antecedent rains.

WATERSHED LAG AND TIME OF CONCENTRATION

The average slope within the watershed together with the overall length and retardance of overland flow are major factors affecting the runoff rate through the watershed.

Time of concentration $(T_{\rm C})$ is the time it takes for water to travel from the most hydraulically distant point in a watershed to its outlet. Lag (L) can be considered as a weighted time of concentration. When runoff from a watershed is nearly uniform it is usually sufficient to relate lag to time of concentration as follows:

$$L = 0.6 T_{c} \tag{5}$$

The lag for the runoff from an increment of excess rainfall can further be considered as the time between the center of mass of the excess

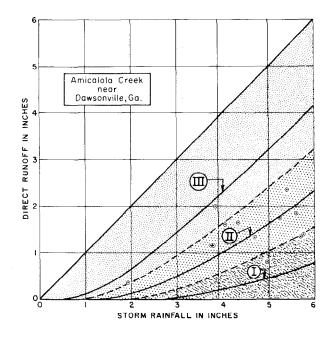
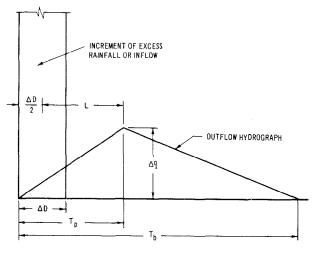


Figure 4.--Limited-gage data used to assign curve numbers to new and unmeasured watershed characteristics.



$$\Delta q = \frac{484 \text{ A} (\Delta Q)}{\frac{\Delta D}{2} + \text{L}} \text{ in C.F.S.}$$

Where:

 $\Delta D = INCREMENT OF STORM PERIOD IN HOURS$

 $\Delta Q = RUNOFF IN INCHES DURING PERIOD <math>\Delta D$

 $\Delta q = PEAK DISCHARGE IN C.F.S. FOR AN INCREMENT OF RUNOFF$

A = DRAINAGE AREA IN SQUARE MILES

 $T_p = TIME TO PEAK (= \frac{\Delta D}{2} + L) IN HOURS$

 $T_b = TIME OF BASE (= 2.67 T_0) IN HOURS$

Figure 5.—Triangular hydrograph relationships.

rainfall increment and the peak of its incremental outflow hydrograph (fig. 5). A graph for estimating lag is shown in figure 6. The equation is:

$$L = \frac{20.8 (s + 1) 0.7}{1900 y^{0.5}}$$
 (6)

Where:

L = lag in hours.

l = length of mainstream to farthest
 divide in feet.

Y = average slope of watershed in percent.

$$S = \frac{1000}{CN^{\frac{1}{2}}} - 10$$

CN' = A retardance factor approximated by
 the curve number representing the
 watershed's hydrologic soil-cover
 complex.

Watershed Shape Factor

The length (1) of the mainstream to the farthest divide was measured on ARS maps of the small experimental watersheds $(\underline{3}, \underline{4}; p. 2.2-7)$

The hydraulic length and area of these watersheds are plotted in figure 7. The relationship is represented by the equation:

$$\ell = 209 \ a^{0.6} \tag{7}$$

Where:

 ℓ = hydraulic length in feet.

a = drainage area in acres.

The ratio of length (ℓ) to average width (w) of a watershed may be referred to as a "shape factor." It follows from equation (7) that the shape factor varies with drainage area.

$$\ell = 43.560 \text{ a/w}$$
 (8)

Where:

w = average width of watershed in feet.

Substituting the value of ℓ in equation (7) for ℓ in equation (8):

$$w = (43,560 \text{ a})/(209 \text{ a}^{0.6})$$

and:

$$w = 208.4 a^{0.4}$$
 (9)

Combining equations (7) and (9):

$$\ell/W = Ka^0 \cdot 2 \tag{10}$$

Where:

 ℓ/w = watershed shape factor.

Variation in shape factor with respect to drainage area based on equation (10) is shown in the following tabulation.

Drainage area (acres)	l/w ¹ Ratio
10	1.58
100	2.51
1000	3.98

 $\underline{1}/$ w is average width of watershed, area/length.

There are small watersheds that do not conform to the shape factor in equation (10); some deviate considerably. In the example shown in figure 8, the diversion terrace along one side changes the shape in reference to the hydraulic length and average width relationship. Here the 1/w factor is 3.75 as compared to a factor of 1.69 based on the general equation (7) used for ES-1026 and ES-1027 solutions. Example 2 under the heading "Basic Procedure for Estimating Peak Discharge Without Developing a Hydrograph" computes the peak discharge for this watershed to be 43 cfs as compared to 46 cfs obtained from the solution in ES-1027. The ES-1026 and ES-1027 solution provides a higher peak discharge estimate for all watersheds that have diversions or terraces and will result in a greater capacity requirement for the design of a structure. This is generally acceptable and often desirable for the installation of smaller measures. Where the economy of a structure requires close adherence to the lesser design capacity, the peak discharge can be determined manually as shown later in example 2. No attempt has been made to modify the precomputed estimates in ES-1026 and ES-1027 for special watershed shape factors since those used change with each change in drainage area as shown by equation (10) and the tabulation following it.

Use of Curye Numbers to Reflect Overland Retardance

The chart for estimating watershed lag in figure 6 uses CN's to reflect the retardance effect of surface conditions on the rate at which runoff moves down the slope. A hay meadow or a thick mulch in a forest is associated with

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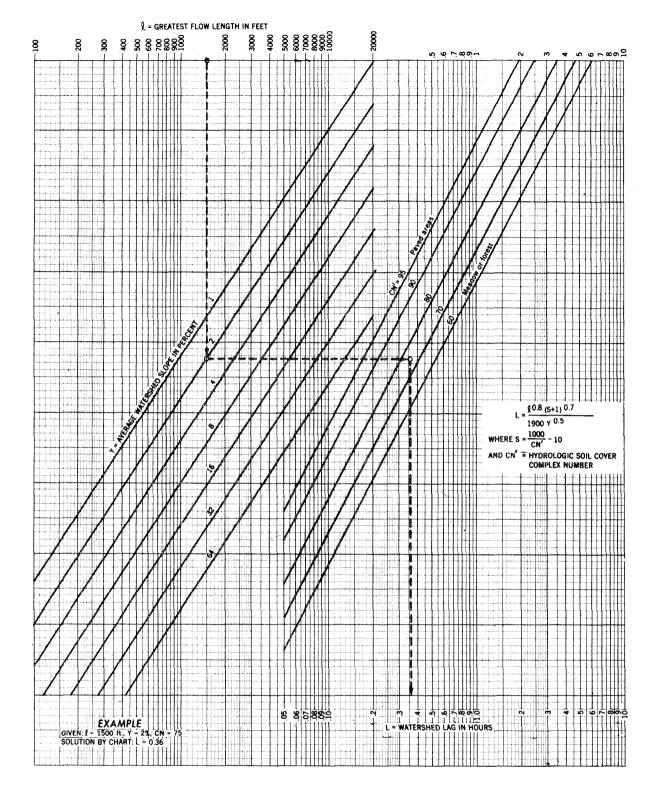


Figure 6.--Watershed lag (NEH-4 January 1971).

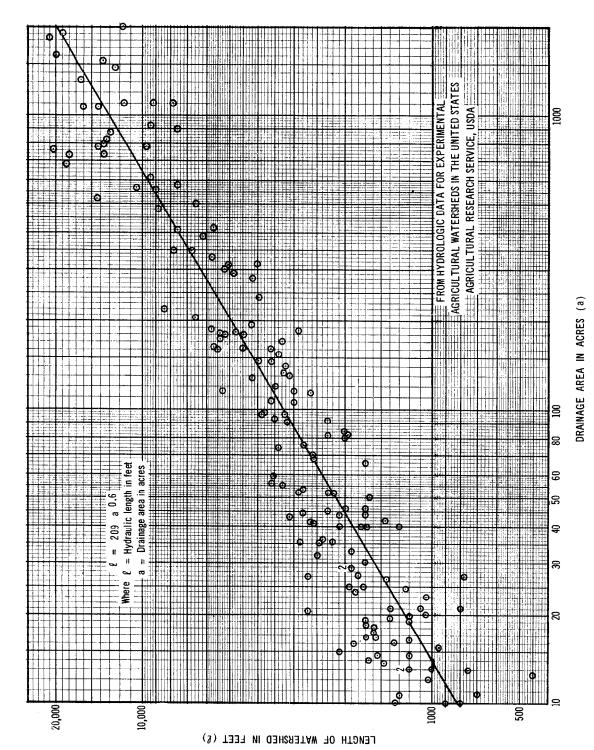


Figure 7.--Hydraulic length and drainage area relationship.

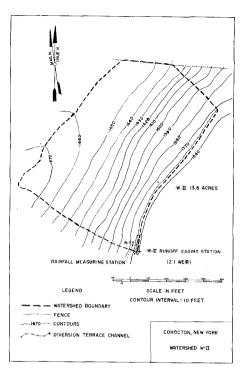


Figure 8.--Natural watershed shape factor altered by a diversion terrace.

low CN's and high retardance. Conversely, a bare surface is associated with high CN's and low retardance. The CN's denoting retardance are the same as those used for estimating the depth of runoff from rainfall (table 2).

The ADP solutions for charts ES-1026 and ES-1027 used the same CN' for computing watershed lag in equation (6) as the CN for depth of runoff in equation (3).

There are unusual situations for which a common CN and CN' does not provide an adequate estimate of peak discharge. One example is a watershed in which the soils have a high infiltration rate (hydrologic soil group A or B) but no surface cover and are on rather steep slopes. Here the CN for estimating depth of runoff is small because of the hydrologic soil group class. Once the soil is saturated and runoff has commenced, however, the overland retardance (CN') for the bare surface is greater than the CN representing the hydrologic soil complex number. In special situations where it is believed that a closer approximation of lag or time of concentration can be made and where a closer peak discharge determination is warranted, the manual solution described later should be made and compared with the results in ES-1026 or ES-1027.

Average Watershed Slope

Slope as used in this method for computing

peak discharge means primarily average watershed slope in the direction of overland flow. Slope is readily available at most locations from existing soil survey data. On larger watersheds the gradient of the stream channel becomes an additional consideration in estimating time of concentration. An estimate of one average slope for all the land within watersheds of less than 2,000 acres is adequate for the slope parameter (Y) in equation (6).

Average slope is defined under three broad categories for the peak discharge charts ES-1026 and ES-1027 (table 4). Peak discharges were computed for the slopes shown in the second column and assigned to the broad categories of the first and third columns. Ordinarily the peak discharge values given for one of the three slope categories in ES-1026 and ES-1027 are adequate for most uses without interpolating between slope categories.

Table 4.--Slope factors for peak discharge computations in charts ES-1026 and ES-1027.

Slope factor	Slope for which computations were made	Average slope range
	Percent	Percent
FLAT ¹ / MODERATE STEEP	1 4 16	0 to 3 3 to 8 8 or more

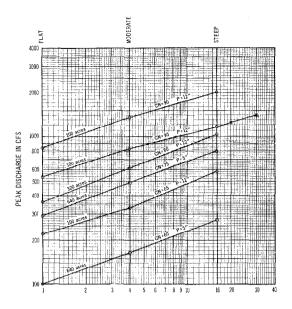
1/ Level to nearly level.

Interpolation for Intermediate Slopes

If a closer estimate of peak discharge is needed than that provided in ES-1026 and ES-1027 for the three slope categories, the value can be determined by interpolation between 1 percent (FLAT), 4 percent (MODERATE), and 16 percent (STEEP). The estimate is made simpler by interpolating along a straight-line plot of peak against slope on log-log paper (fig. 9). The straight-line plot on log-log paper can also be used to extrapolate peak discharge values for slopes steeper than 16 percent. But other parameters than those in equation (6) may need to be considered for average watershed slopes steeper than 30 percent.

TRIANGULAR HYDROGRAPH EQUATION

The triangular hydrograph is a practical representation of excess runoff with only one rise, one peak, and one recession. It has been



AVERAGE WATERSHED SLOPE IN PERCENT

Figure 9.--Logarithmic interpolation of peak discharge for intermediate slopes.

very useful in the design of soil and water conservation measures. Its geometric makeup can be easily described mathematically, which makes it very useful in the processes of estimating discharge rates.

SCS developed the following equation to estimate the peak rate of discharge for spillway and channel capacities for conservation measures and water-control structures:

$$q_p = (KAQ)/T_p$$
 (11) (2, 5, 7)

Where:

 q_{p} = peak rate of discharge.

A = drainage area contributing to the peak rate.

Q = storm runoff.

K = a constant.

 $T_D = time to peak.$

Time to peak (T_p) is expressed as:

$$T_p = \frac{D}{2} + L$$

Where:

D = storm duration.

INCREMENTAL HYDROGRAPHS

Total storm rainfall rarely if ever occurs uniformly with respect to time. Because rainfall gage data and the variation of rainfall with time are lacking in most small watersheds, it is desirable that variations in rainfall with respect to time be standardized for the design of soil and water conservation measures. To use equation (11) for other than uniform storm rainfall, it is necessary to divide the storm into increments of duration (AD) and compute corresponding increments of runoff (AQ) (fig. 5). The peak discharge equation for an increment of runoff is:

$$\Delta q_{p} = \frac{484A (\Delta Q)}{\frac{\Delta D}{2} + L}$$
 (12)

Where:

A is in square miles.

 ΔQ is in inches.

AD and L are in hours.

 Δq_p is in cfs.

The constant, K, in equation (11) becomes 484 when the peak discharge is computed in units of cfs for the triangular hydrograph (fig. 5). The ordinates of the individual triangular hydrographs for each Δq_p are added to develop a composite hydrograph (fig. 10). Note how each incremental hydrograph is displaced one ΔD to the right for each succeeding time increment.

BASIC PROCEDURE FOR ESTIMATING PEAK DISCHARGE WITHOUT DEVELOPING A HYDROGRAPH

The plotting and summation of unit hydrograph ordinates (fig. 10) require more time than desirable or necessary to obtain only the peak discharge (qp) for a design storm. The peak discharge, without the further development of the entire composite hydrograph, is all that is required for most SCS applications. For these the solution can be reduced to the period of runoff or of excess rainfall that directly affects the peak rate corresponding to a given watershed lag (L). A relationship between AD and L can be chosen that enables the summation of only a single ordinate from each incremental hydrograph within the effective runoff period to compute the peak discharge. The usual choice is to make ΔD equal to one—third the time to peak (Tp) (fig. 11). If ΔD is taken to equal Tp/3 then the equation for AD is:

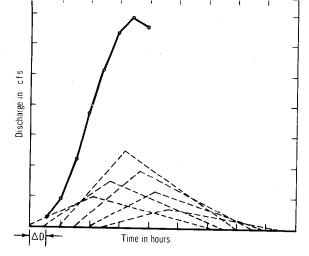


Figure 10.--Composite hydrograph from hydrographs for storm increments ΔD_{\bullet}

$$\Delta D = 0.4L \tag{13}$$

Where:

$$T_p = (\Delta D/2) + L \text{ (fig. 5)}$$

and

$$T_p = 3 \Delta D$$

The effective peak-producing runoff period is $7\Delta D$ with the fifth increment ΔD_5 being the most intense runoff increment (fig. 12). The peak discharge for each increment (Δq_p) can be computed by equation (12) using:

$$\Delta Q_1 = \text{Mass } Q_2 - \text{Mass } Q_1$$

$$\Delta Q_2 = \text{Mass } Q_3 - \text{Mass } Q_2 \text{ etc.} \tag{14}$$

SELECT
$$\Delta D = 1/3 T_p$$
 OR $T_p = 3 \Delta D$

SINCE
$$T_p = \frac{\Delta D}{2} + L$$
, $\Delta D = 0.4L$

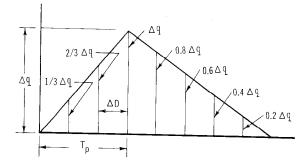


Figure 11.--Making ΔD equal to one-third the time to peak.

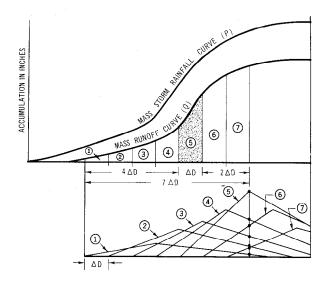


Figure 12.--Effective peak-producing period and most effective increment.

The y values in figure 13 are the proportional parts of Δq_{p1} , Δq_{p2} , etc., contributing to the composite peak q_p , Δq_p having been obtained for each by equation (12). The product $(y)\Delta q_p$ for each of the seven increments of runoff are added to obtain the composite peak rate (q_p) . The summation equation is:

$$q = \Sigma \cdot 0.2\Delta q_1 + 0.4\Delta q_2 + 0.6\Delta q_3 + 0.8\Delta q_4 + 1.0\Delta q_5 + \frac{2}{3}\Delta q_6 + \frac{1}{3}\Delta q_7$$
 (15)

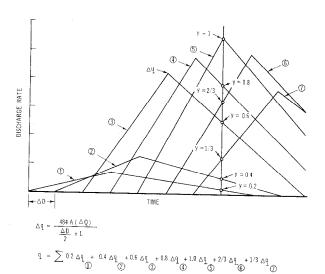


Figure 13.--Proportional parts of incremental hydrographs that contribute to the composite peak.

The equations were solved by ADP to get the peak-discharge rates for ES-1026 and ES-1027. These equations can be solved manually by following the examples given here.

Example 1.-- Given a 100-acre watershed with runoff characteristics represented by CN 80 in table 2. The average slope of the watershed is 1 percent. The peak discharge is required for a 10-inch rain in 2^{14} hours. The watershed is located in the area covered by the type II curve in figure 1.

Step 1.--Estimate the hydraulic length of the watershed by equation (7):

 $l = 209a^{0.6}$ $l = 209(100)^{0.6}$ l = 3,300 feet

Step 2.—Read watershed lag from figure 6 for l = 3.300 feet; Y = 1 percent and CN 80:

L = 0.83 hour

Step 3.--Compute ΔD from equation (13), assuming ΔD = $T_{\rm D}/3$:

 $\Delta D = 0.4L$ $\Delta D = 0.4(0.83)$ $\Delta D = 0.33 \text{ hour}$

Step 4.--Compute the effective peak-producing runoff period for $7\Delta D$:

 $7\Delta D = 7(0.33)$ hour $7\Delta D = 2.31$ hours

Step 5.--Prepare a tabulation based on a type II distribution in table 1; P_{24} = 10 inches and runoff (Q) for CN 80 from figure 3:

Time (hours)	P _X /P ₂₄	Mass P (inches)	Mass Q (inches)
10.0 10.5 11.0 11.5 11.75 12.0 12.5 13.0	0.181 .204 .235 .283 .387 .663 .735	1.81 2.04 2.35 2.83 3.87 6.63 7.35 7.72	0.44 •59 •78 1.12 1.94 4.36 5.02 5.36

Step 6.—Prepare working curve. Plot mass Q versus time (fig. 14). Select midpoint of maximum increment of runoff (11.88 hours). This will be the same for most type II distributions, but it will occur later where initial abstraction ($I_a = 0.28$) has not been satisfied prior to 11.75 hours. Mark the curve with the 7 Δ D beginning at 10.39 hours for the selected midpoint minus 4.5 Δ D.

$$11.88 - 4.5(0.33) = 10.39$$

Step 7.--Prepare computations for instantaneous peak discharge (table 5). The increment in

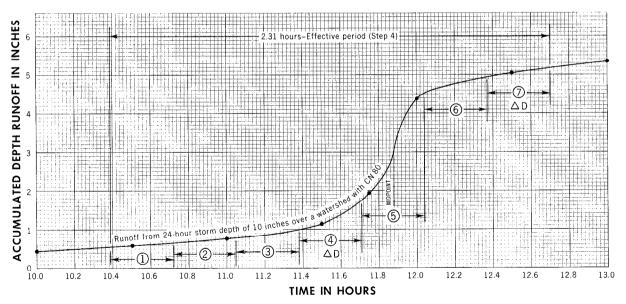


Figure 14. -- Working curve for manual computation from type II storm distribution, table 1.

Table 5.--Example 1, computations for instantaneous peak discharge

(1)	(2)	(3)	(4)	(5)	(6)	(7)
Increment	Time	Mass runoff	∆ଢ଼	$\Delta q^{\frac{1}{2}}$	<u>y2</u> /	Y(Aq)
	Hours	Inches	Inches	Cfs		Cfs
	10.39	0.55	0.10	0.3	0.0	- 0
ΔD_1	10.72	.67	0.12	9.1	0.2	1.8
ΔD_2			.13	9.9	•)1	4.0
ΔD_{3}	11.05	.80	.18	13.7	.6	8.2
3	11.38	.98	•77	58.5	.8	46.8
$\Delta D_{t_{\dagger}}$	11.71	1.75	• 1 [JU•J	•0	40.0
ΔD_{5}	12.04	4.53	2.78	211.3	1.0	211.3
ΔD_{6}	12.04	4.73	.42	31.9	2/3	21.3
	12.37	4.95	,22	16.7	1/3	5.6
ΔD,7	12.70	5.17	• ८८	70.1	-17	·
						299.03/

 $[\]frac{1}{r}$, From equation (12) $\Delta q = 76.0 (\Delta Q)$

 $\frac{2}{}$ See figure 13

column 1 and the time in column 2 correspond with the beginning and end of each incremental period, ΔD , in figure 14. The runoff (Q) in column 3 is read from the curve in figure 14. Column 4 is the incremental runoff for each ΔD . Peak discharge for each increment of runoff is computed by equation (12) and tabulated in column 5. Column 6 lists the proportion of incremental peak that contributes to the total peak as shown in figure 13. Column 7 is the summation of proportionate parts of each incremental peak in equation (15).

Example 2.—Given watershed W-II, 13.8 acres located at Cohocton, N. Y. The watershed is in cultivation with good conservation treatment in effect; its soils are predominantly in hydrologic soil group C. The average watershed slope is 20 percent and hydraulic length & is measured as 1,500 feet following the course of the diversion terrace (fig. 8). The peak discharge for a 25-year frequency storm is desired for AMC II.

Step 1.--Select CN from table 2 based on the watershed description: CN = 82

Step 2. -- Compute S from equation (4):

$$S = \frac{1000}{CN} - 10$$

$$S = \frac{1000}{82} - 10$$

Step 3.—Read watershed lag (L) from figure 6 or compute L from equation (6):

$$L = 0.1 \text{ (approx.)}$$

Step 4.--The 24-hour, 25-year frequency rainfall for Cohocton, N. Y., in the Weather Bureau Atlas is 4.3 inches. Use type II distribution. Step 5.--Compute ΔD from equation (13) assuming ΔD = $T_{\rm p}/3$:

$$\Delta D = 0.4L$$

$$\Delta D = 0.4(0.1) = 0.04 \text{ hour}$$

Step 6.--Compute the effective peak-producing runoff period for $7\Delta D$:

$$7\Delta D = 7(0.04) \text{ hour}$$

$$7\Delta D = 0.28 \text{ hour}$$

Step 7.--Prepare a tabulation from data in steps 1 and $^{\rm h}$ for the period in step 6, solving for Q by using equation (3) or by reading Q from figure 3:

$$P = 4.3$$
 inches; $S = 2.2$ inches.

Time (hours)	$P_{x}/P_{24}\frac{1}{}$	Mass P (inches)	Mass Q (inches)
11.5	0.283	1.22	0.20
11.75	.387	1.66	.44
12.0	.663	2.85	1.26

^{1/}From table 1, type II distribution.

Step 8.—Prepare working curve (fig. 15) from data in step 7.

Step 9.--Prepare computations for instantaneous peak discharge (table 6).

The peak discharge for this example is rounded to $\frac{13}{43}$ cfs, as computed manually, and by estimating $\overline{\text{lag}}$ (L) on the actual hydraulic length (l) along the diversion terrace. The peak discharge obtained from ES-1027 (sheets 19 and 20), with l based on equation (7) and not the measured length along the diversion terrace, is:

- q for STEEP, CN 80, 13.8 agres, and P = 4.3 inches is 43 cfs.
- q for STEEP, CN 85, 13.8 acres, and P = 4.3 inches is 50 cfs.

By interpolation,

q for STEEP, CN 82, 13.8 acres, and P = 4.3 inches is 46 cfs.

Converting from the 16-percent slope for STEEP to a 20-percent slope would not add more than 1

 $^{^{2}/}_{\rm qp}$ = 300 (approx) from ES-1027, Rev. 2-15-71 sheet 5 of 21.

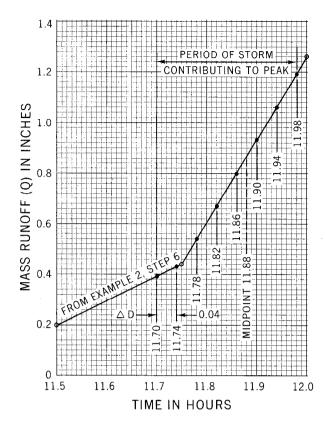


Figure 15. -- Working curve for example 2.

or 2 cfs by extrapolation on log-log paper as was suggested for special cases (fig. 9).

It may be concluded that the ES-1027 charts overestimate the peak discharge in this example by about 3 cfs or 7 percent. This is due mainly to the alteration of the watershed shape factor by the diversion terrace.

Example 3.--This example demonstrates the need for making ΔD smaller than 0.4L as used in the previous two examples. To keep it less than 0.5 hour and more commensurate with the increment of maximum storm intensity in table 1, it is set equal to 1/6 T_p instead of 1/3 T_p and it follows that:

$$\Delta D = 0.182L$$
 (16)

Given a 2,000-acre watershed with CN 60 and an average slope of 8 percent located on Kahuka Point, Oahua, Hawaii. An estimate of the peak discharge for a 25-year frequency rainstorm is desired.

Step 1.--Estimate the hydraulic length of the watershed by equation (7) or read from figure 7:

Table 6.--Example 2, computations for instantaneous peak discharge

(1)	(2)	(3)	(4)	(5)	(6)	(7)
Increment	Time Hours	Mass runoff Inches	ΔQ Inches	Δq½/ Cfs	Y	Υ(Δq) Cfs
4.50	11.702/	0.39	0.04	3.5	0.2	.7
ΔD_1	11.743/	0.43	0.04	9.6	0.2	3,8
ΔD_3	11.78	0.54	0.13	11.3	0.6	6.8
ΔD,	11.82	0.67	0.13	11.3	0.8	9.0
ΔD_5	11.86	0.80	0.13	11.3	1.0	11.3
ΔD ₆	11.90	0.93	0.13	11.3	2/3	7.5
ΔD_{7}	11.98	1.19	0.13	11.3	1/3	3.8
				TOT	AL =	42.9

$$\frac{1}{\Delta Q} = \frac{484 \text{ A } (\Delta Q)}{\frac{\Delta D}{Q} + L} = \frac{(484) (13.8) (\Delta Q)}{(0.02 + 0.1) 640} = 87.0 \text{ AQ}$$

$$\frac{2}{11.88} - 4.5 \Delta D = 11.88 - 4.5(0.4) = 11.70$$

$$\frac{3}{11.70} + \Delta D = 11.70 + 0.04 = 11.74$$
 hours (etc.)

Step 2.—Read watershed lag from figure 6 for $\ell = 20,000$ feet; Y = 8 percent and CN' 60:

$$L = 2.1 \text{ hours}$$

Step 3.--Compute ΔD from equation (16), assuming ΔD = $T_{\rm p}/6$:

$$\Delta D = 0.38 \text{ hour}$$

Step 4.—Compute the effective peak-producing runoff period for $15\Delta D$:

$$15\Delta D = 15(0.38)$$
 hour

$$15\Delta D = 5.7$$
 hours

Step 5.--Prepare a tabulation based on a type I distribution in table 1; P_{24} = 10 inches and CN 60.

Time	$P_{\rm X}/P_{\rm 24}$	Mass P	Mass Q
(hours)		(inches)	(inches)
6.00 7.00 8.00 8.50 9.00 9.50 9.75 10.00 10.50 11.00 11.50 12.00	0.125 .156 .194 .219 .254 .303 .362 .515 .583 .624 .654	1.25 1.56 1.94 2.19 2.54 3.03 3.62 5.15 5.83 6.54 6.54	0.00 .00 .05 .10 .18 .35 .59 1.39 1.82 2.08 2.28

Step 6.--Prepare working curve (fig. 16) from data in step 5.

Step 7.--Prepare computations for instantaneous peak discharge (table 7).

EQUATIONS AND ASSUMPTIONS USED IN COMPUTER SOLUTIONS FOR CHARTS ES-1026 AND ES-1027

Storm Rainfall

Fifteen- and 30-minute increments of accumulated-to-total ratios of rainfall were used with both type I and II distributions shown in figure 1. The 15-minute increments extended through the most intense 1-hour period of each distribution. Twenty-four-hour storms were generated accordingly for each distribution for those rainfall depths shown in the ES charts.

Rainfall-Runoff Equations

Runoff (Q) was computed accumulatively from the two accumulated rainfall distributions and their increments described. This solution was made for all rainfall depths and for each of the seven CN's included in the ES charts by the following equations:

$$Q = \frac{(P - 0.2S)^2}{P + 0.8S} \tag{3}$$

and

$$S = \frac{1000}{CN} - 10 \tag{17}$$

Watershed Lag

Lag time (L) was computed for 1-, 4-, and 16percent slopes (Y) for each of the seven CN's in the ES charts and for each of the following drainage areas (a):

5 acres

10 to 100 acres by 10-acre increments 100 to 1,000 acres by 20-acre increments 1,000 to 2,000 acres by 50-acre increments

The programmed equations were:

$$L = \frac{10.8 (s + 1)^{0.7}}{1900 y^{0.5}}$$
 (6).

$$l = 209 \ a^{0.6}$$
 (7)

$$S = \frac{1000}{CN} - 10 \tag{17}$$

CN' for computing I, is approximated by the CN from table 2.

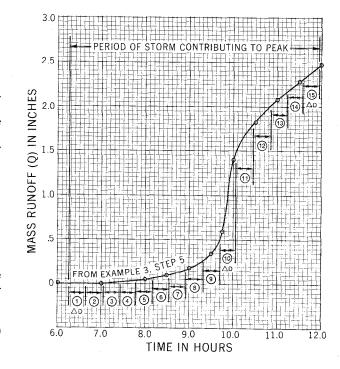


Figure 16.--Working curve for example 3.

Period of Runoff Affecting Peak Discharge

The computer program related the incremented periods (ΔD) of storm runoff to lag (L) as in (example 3):

$$\Delta D = 0.182 L$$
 (16)

The peak producing storm period for this relationship is 15 ΔD (table 7, example 3).

The computer solution determined the time at which the midperiod of the most intense 15-minute increment of accumulated runoff occurred. This was at 9.875 hours for the type I distribution and 11.875 hours for the type II distribution. It computed the time at the beginning of the effective period (15AD) as:

Incremental Peak Discharge

The instantaneous peak discharge was computed for each increment of runoff (ΔQ) within the

Table 7.--Example 3, computations for instantaneous peak discharge

(1)	(2)	(3) Mass	(4)	(5)	(6)	(7)
Increment	Time	runoff	ΔQ	<u>Δq</u> 1/	Y	$Y(\Delta_q)$
	Hours	Inches	Inches	Cfs		Cfs
AD	6.27 ² /	0.00	0.00	0	0.1	0
ΔD_1	6.65 <u>3</u> /	.00				
ΔD_2	7.03	.00	.00	0	.2	0
ΔD ₃			.00	0	•3	0
$\Delta D_{t_{\bullet}}$	7.41	.02	.02	13	.4	5
ΔD ₅	7.79	.04	.03	20	•5	10
ΔD ₆	8.17	.07	.04	26	.6	16
Ť	8.55	.11				
ΔD ₇	8.93	.17	.06	40	•7	28
ΔD 8	9.31	.26	.09	59	.8	47
$_{ m e}$ ${ m d}\Delta$.23	152	•9	137
AD ₁₀	9.69	.49	1.00	660	1.0	660
ΔD 11	10.07	1.49	.31	205	5/6	171
	10.45	1.80	.20	132	4/6	88
ΔD ₁₂	10.83	2.00				
ΔD ₁₃	11.21	2.17	.17	112	3/6	56
ΔD _{1 4}	11.59	2.32	.15	99	2/6	33
ΔD ₁₅			.13	86	1/6	14
	11.97	2.45	me m			1 <u>265</u> cfs
		_	1017	AL q =		1205 cfs

 $[\]frac{1}{2}$ $\Delta q = \frac{\frac{484}{\Delta D}}{\frac{\Delta D}{2}} + L (\Delta Q) = \frac{\frac{484}{2.29}}{(640)} (\Delta Q) = 660 (\Delta Q)$

effective period (15 Δ D) described according to the following equation:

$$\Delta q = \frac{0.69a}{L}(\Delta Q) \tag{18}$$

Combined Peak Discharge

The incremental peaks (Δq 's) were combined in the computer program in a manner similar to the manual solution shown in table 7, example 3.

 $[\]frac{2}{9.88} - 9.5\Delta D = 9.88 - 9.5(.38) = 6.27$

 $[\]frac{3}{6.27} + \Delta D = 6.27 + .38 = 6.65 \text{ hours(etc.)}$

- (1) Chow, Ven te. 1962. Hydrologic determinations of waterway areas for the design of drainage structures in small drainage basins. Ill. Engr. Expt. Sta. Bull. 462, 104 p.
- (2) Ogrosky, Harold O., and Victor Mockus. 1964. Hydrology of agricultural lands. <u>In</u> Handbook of applied hydrology, Ven te Chow, ed., (sec. 21), 97 p. McGraw-Hill Book Co., New York.
- (3) U.S. Agricultural Research Service. 1963. Hydrologic data for experimental agricultural watersheds in the United States 1956-59. Misc. Publ. 945. 611 p.

(4) 1960. Selected runoff events for small agricultural watersheds in the United States. 374 p.

(5) U.S. Bureau of Reclamation. 1960. Design of small dams (appendix A). 611 p.

(6) U.S. National Weather Service. 1973. Precipitation-frequency atlas of western United States. NOAA atlas No. 2, v. 1-11.

(7) U.S. Soil Conservation Service. 1972. Hydrology. Nat. Eng. Handb., sec. 4. 547 p. (8) U.S. Weather Bureau. 1963. Probable maximum precipitation and rainfall-frequency data for Alaska for areas to 400 square miles, durations to 24 hours, and return periods from 1 to 100 years. Tech. Paper 47. 69 p.

(9) 1962. Rainfall-frequency atlas for the Hawaiian Islands for areas to 200 square miles, durations to 24 hours, and return periods from 1 to 100 years. Tech.

Paper 43.

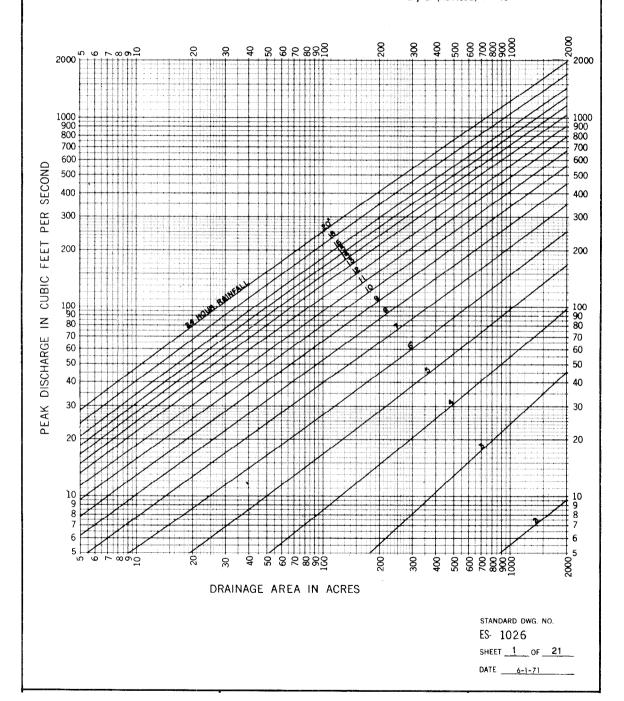
(10)

1961. Generalized estimate of probable maximum precipitation and rainfall-frequency data for Puerto Rico and Virgin Islands for areas to 400 square miles, durations to 24 hours, and return periods from 1 to 100 years. Tech. Paper 42. 94 p.

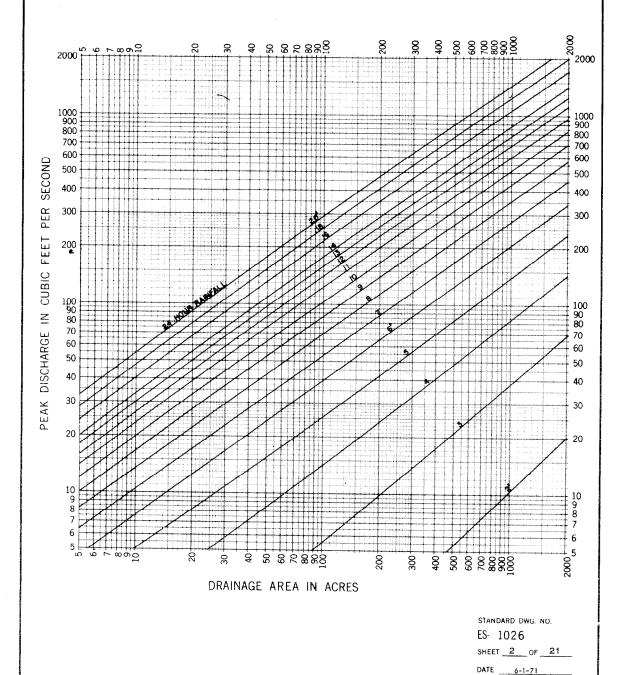
(11) 1961. Rainfall-frequency atlas of the United States for durations from 30 minutes to 24 hours and return periods from 1 to 100 years. Tech. Paper 40.

115 p.

SLOPES - FLAT CURVE NUMBER - 60



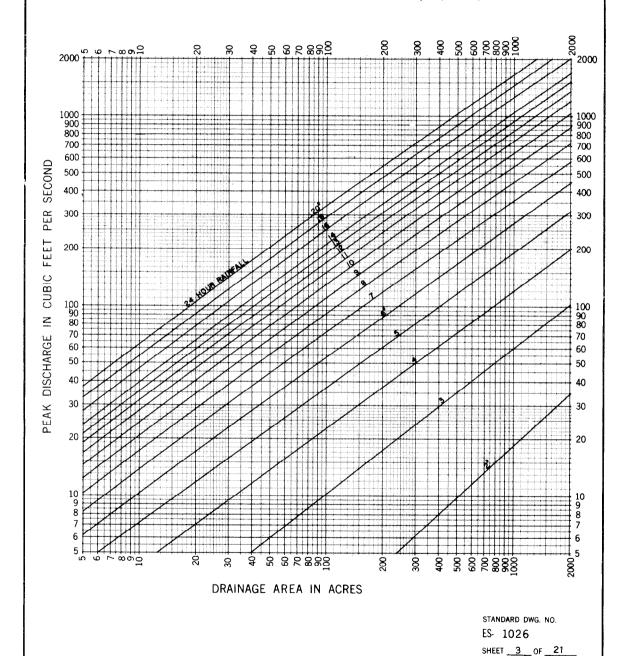
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SLOPES - FLAT

CURVE NUMBER - 70

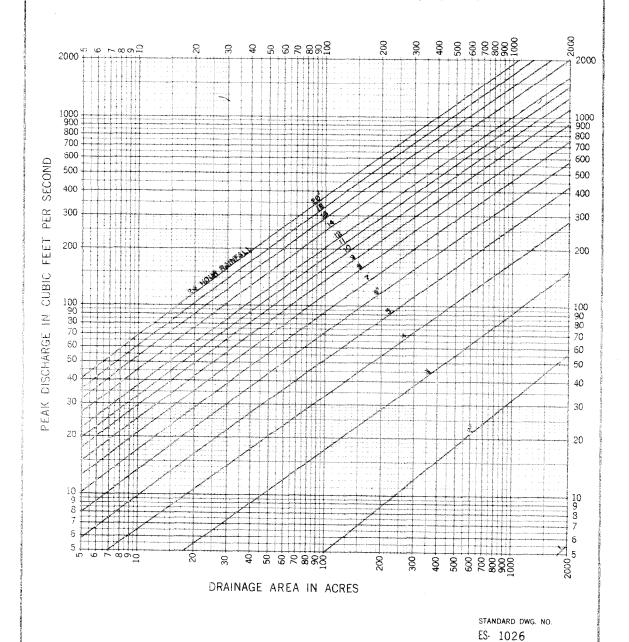
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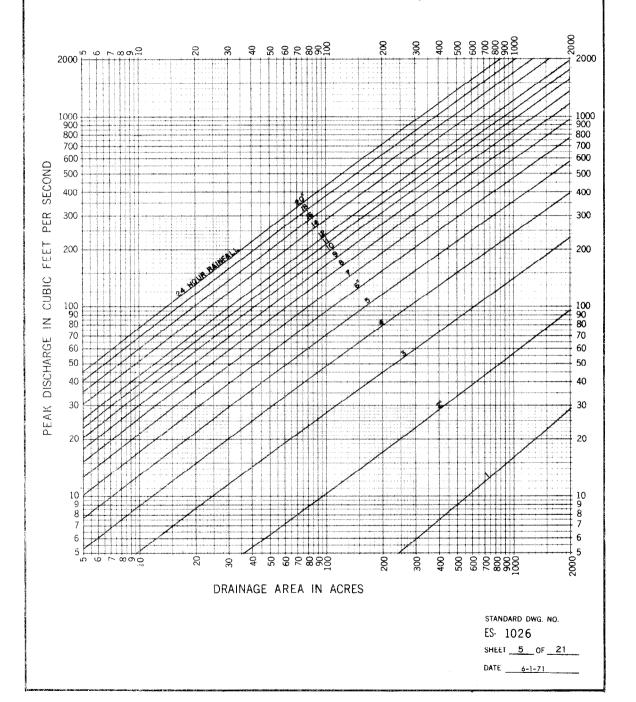
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SHEET 4 OF 21

DATE 6-1-71

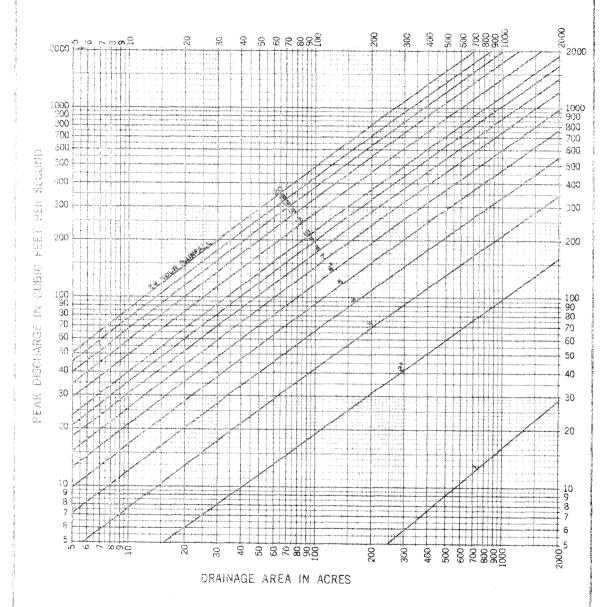


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SLOPES - FLAT CURVE NUMBER - 85

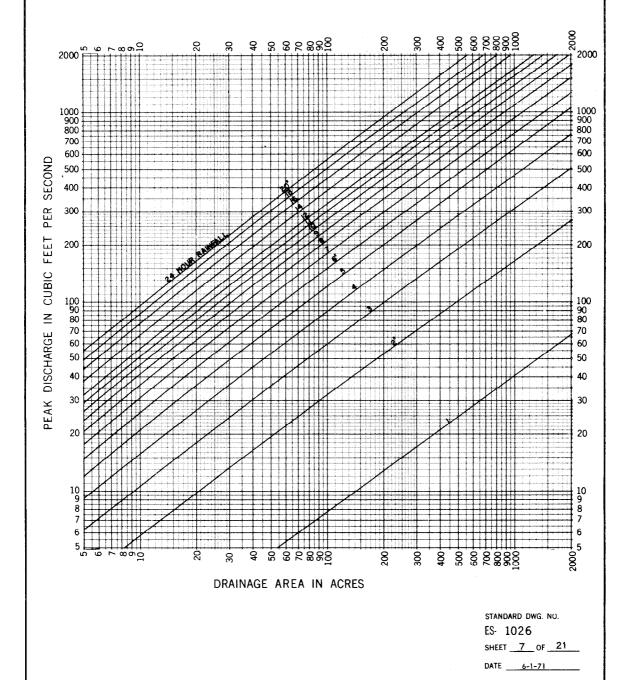
24 HOUR RAINFALL FROM US WB TP-43, TP-47, & (Revised) TP-40



STANDARD DWG. NO.
ES- 1026
SHEET <u>6</u> OF <u>21</u>

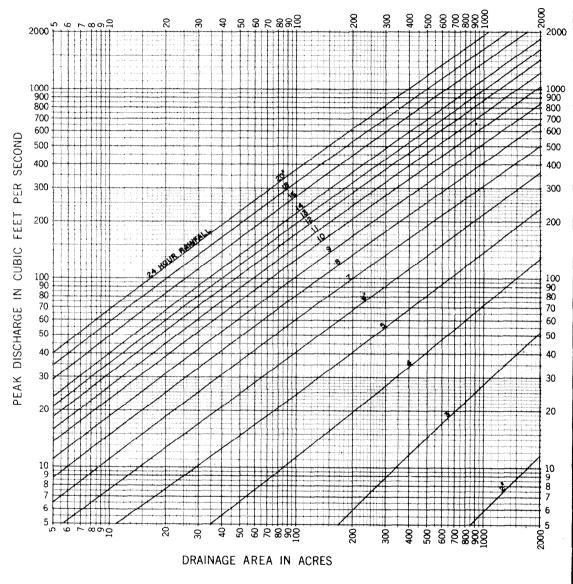
DATE 6-1-71

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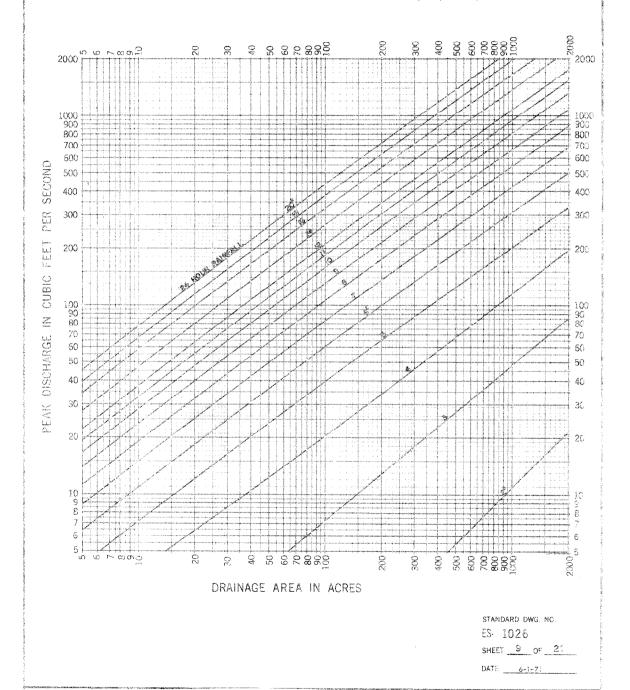
24 HOUR RAINFALL FROM US WB TP-43, TP-47, & (Revised) TP-40



STANDARD DWG. NO.
ES- 1026
SHEET <u>8</u> OF <u>21</u>

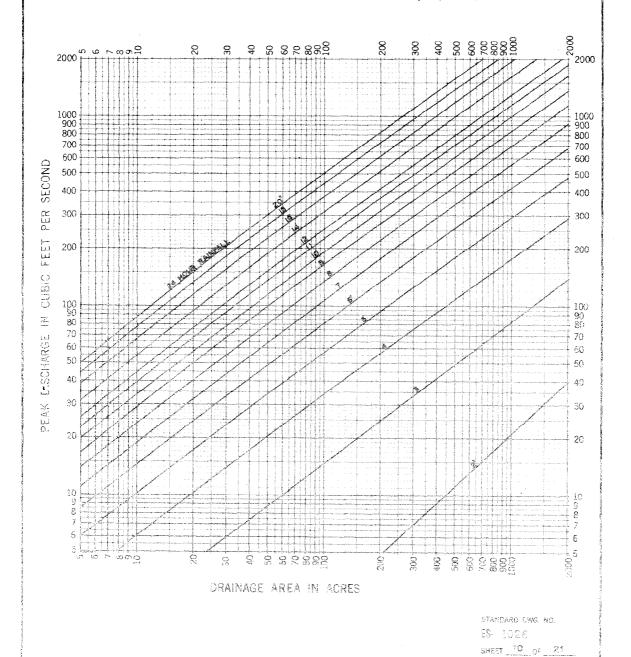
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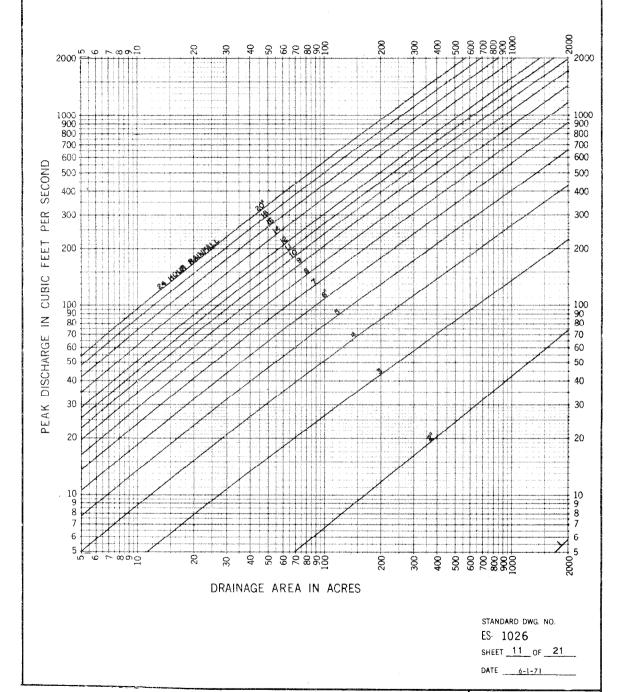


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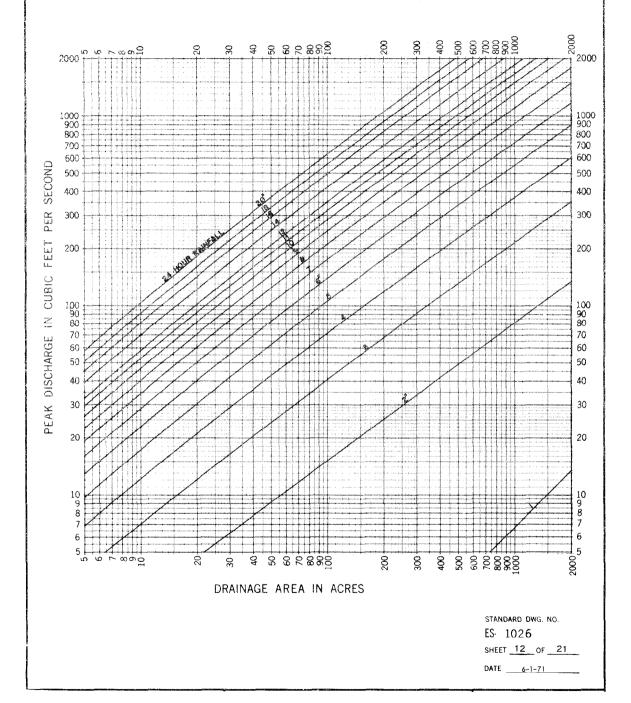
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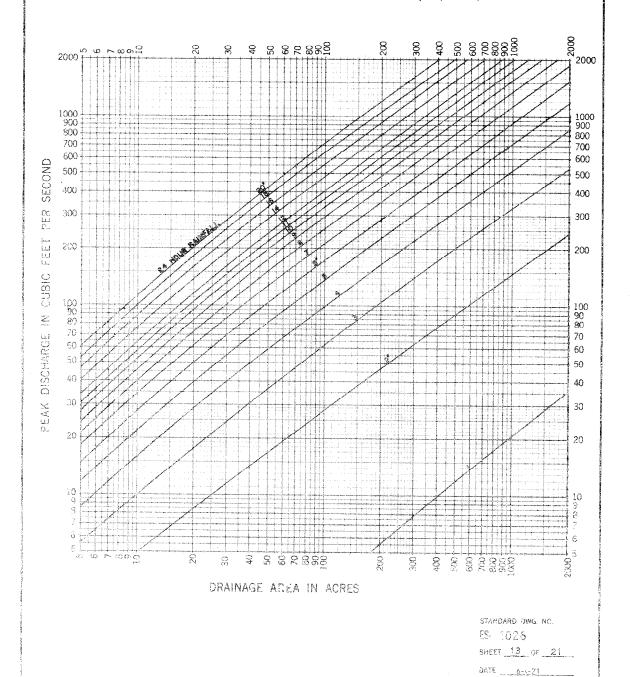
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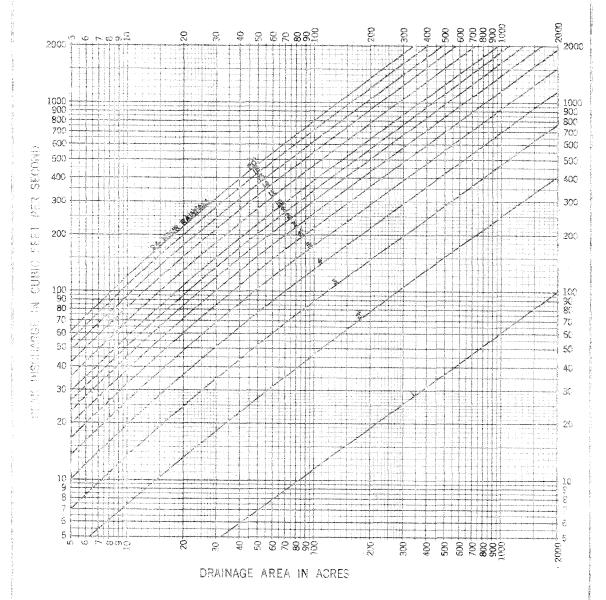


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SLOPES - MODERATE
CURVE NUMBER - 90

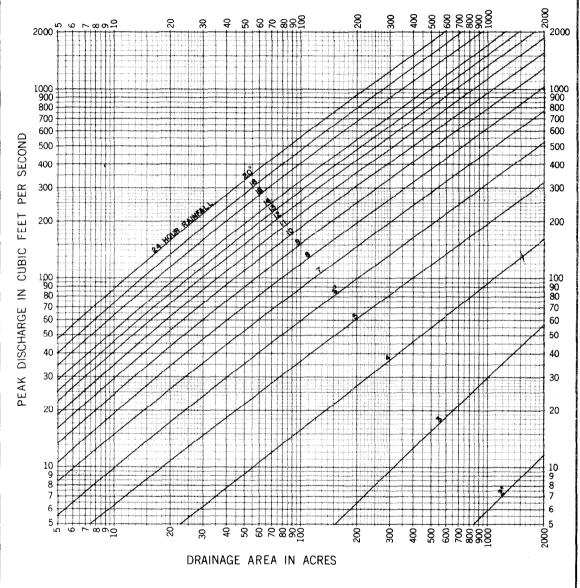
24 HOUR RAINFALL FROM US WB TP-43, TP-47, & (Revised) TP-40



STANDARD DWG. NO.
ES- 1026
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DATE 6-1-7:

SLOPES - STEEP CURVE NUMBER - 60

24 HOUR RAINFALL FROM US WB TP-43, TP-47, & (Revised) TP-40



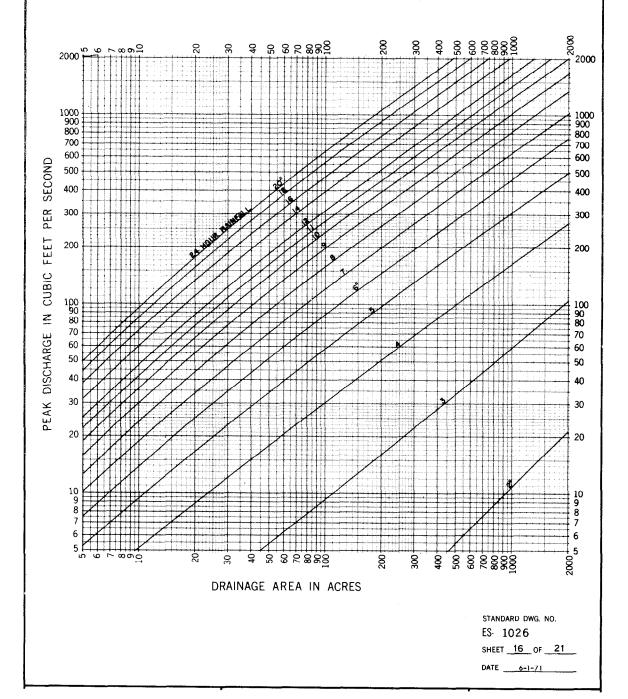
STANDARD DWG. NO. ES- 1026

SHEET 15 OF 21

DATE 6-1-71

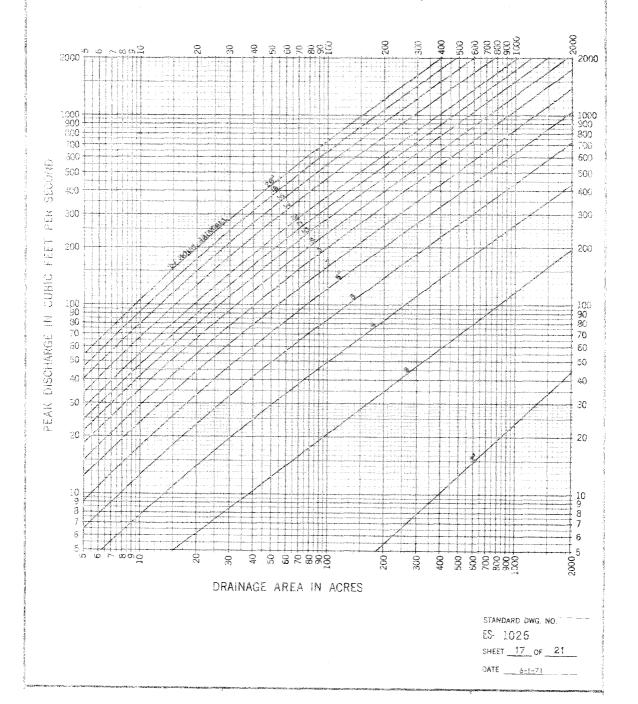
SLOPES - STEEP CURVE NUMBER - 65

24 HOUR RAINFALL FROM US WB TP-43, TP-47, & (Revised) TP-40



SLOPES - STEEP CURVE NUMBER - 70

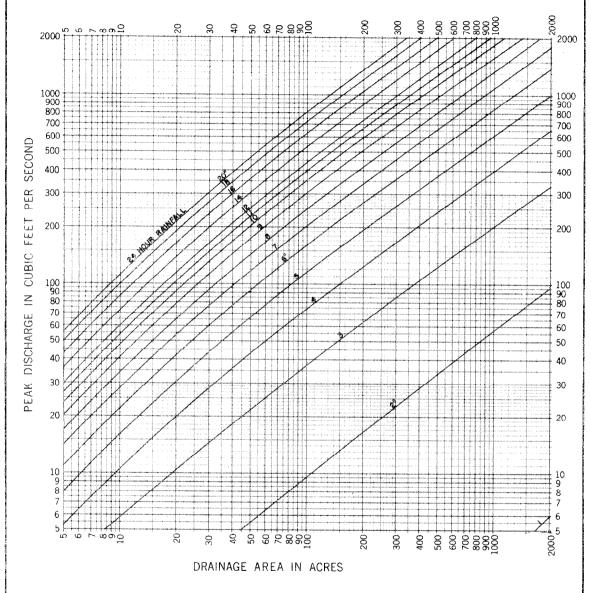
24 HOUR RAINFALL FROM US WB TP-43, TP-47, & (Revised) TP-40



SLOPES - STEEP

CURVE NUMBER - 75

24 HOUR RAINFALL FROM US WB TP-43, TP-47, & (Revised) TP-40



STANDARD DWG. NO. ES- 1026

SHEET 18 OF 21

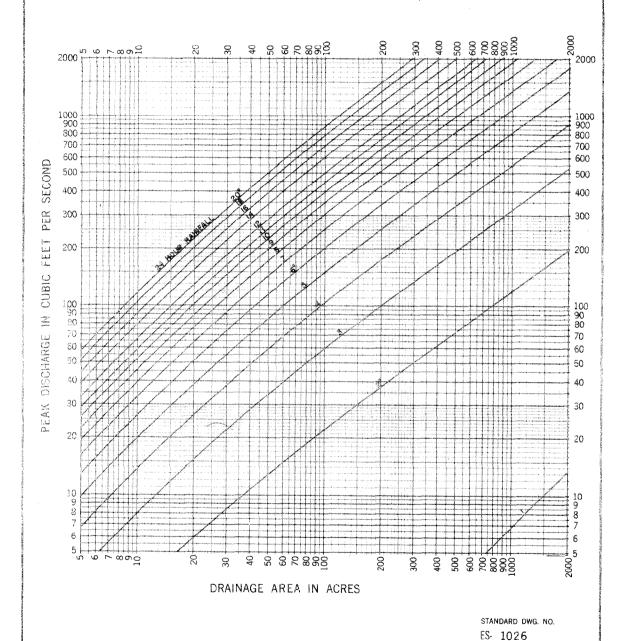
DATE 6-1-71

SLOPES - STEEP CURVE NUMBER - 80

SHEET 19 OF 21

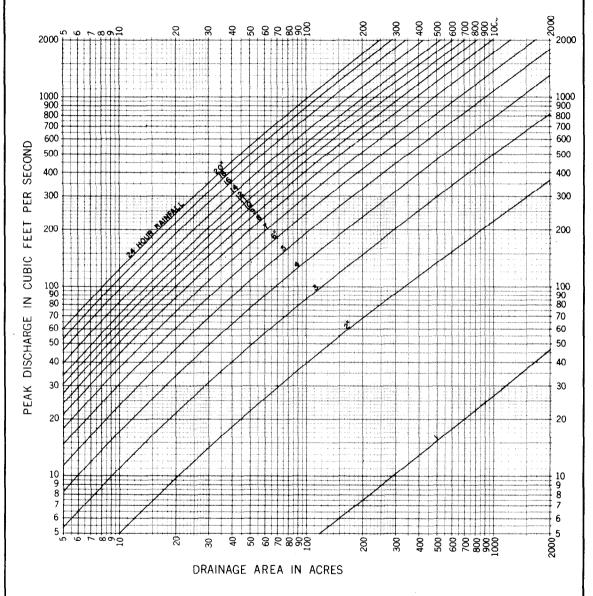
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24 HOUR RAINFALL FROM US W8 TP-43, TP-47, & (Revised) TP-40



SLOPES - STEEP CURVE NUMBER - 85

24 HOUR RAINFALL FROM US WB TP-43. TP-47, & (Revised) TP-40

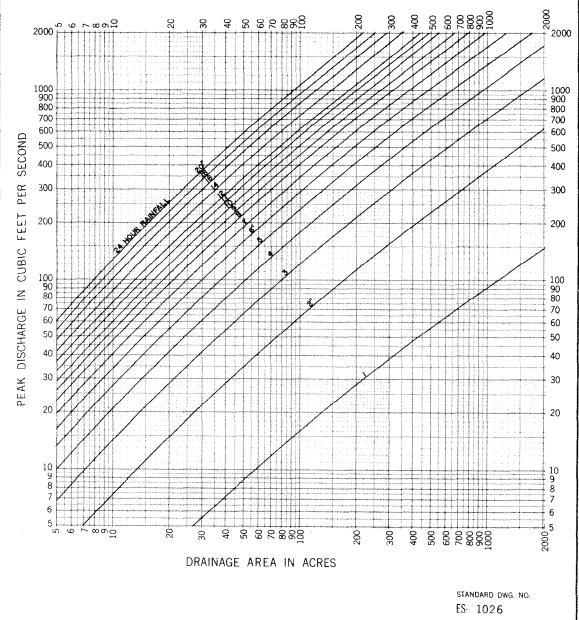


STANDARD DWG. NO. ES- 1026 SHEET 20 OF 21

DATE 6-1-71

SLOPES - STEEP CURVE NUMBER - 90

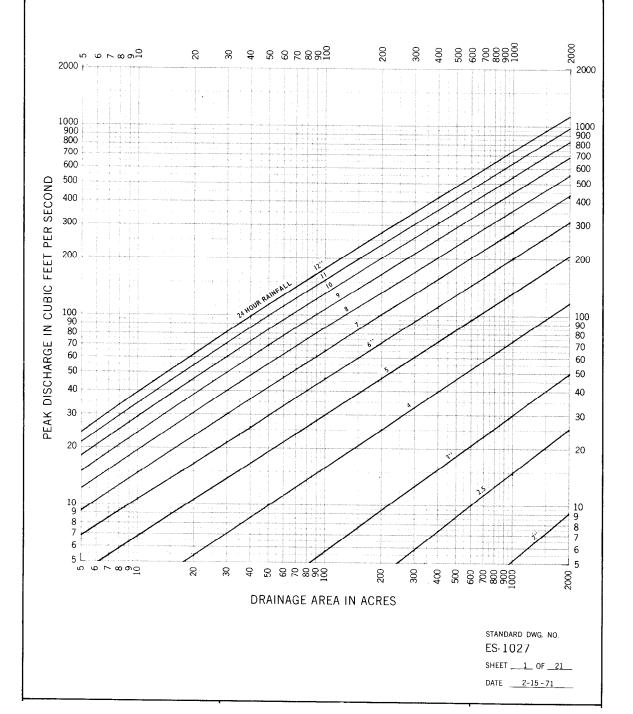
24 HOUR RAINFALL FROM US WB TP-43, TP-47, & (Revised) TP-40



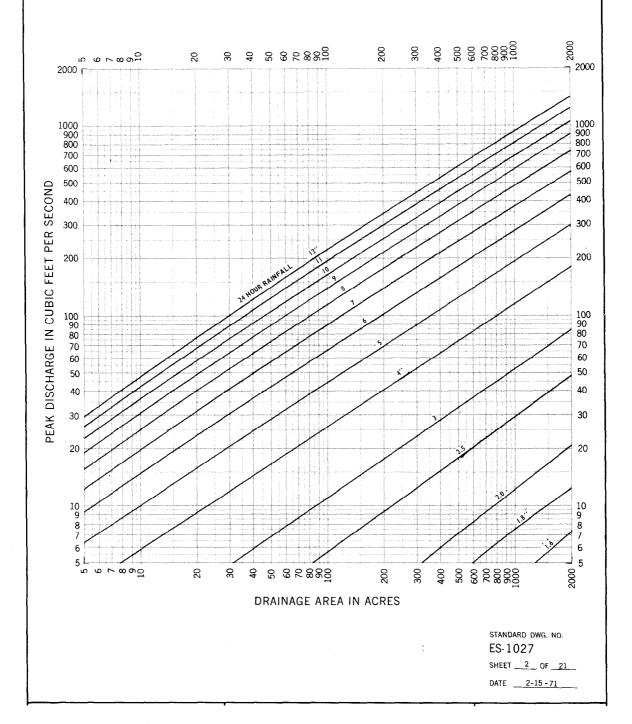
SHEET 21 OF 21

DATE ____6-1-71

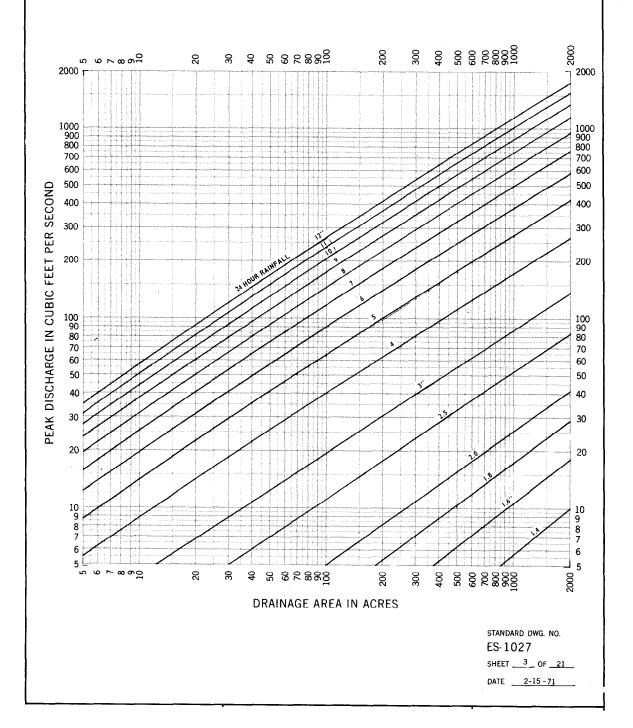
SLOPES - FLAT CURVE NUMBER - 60



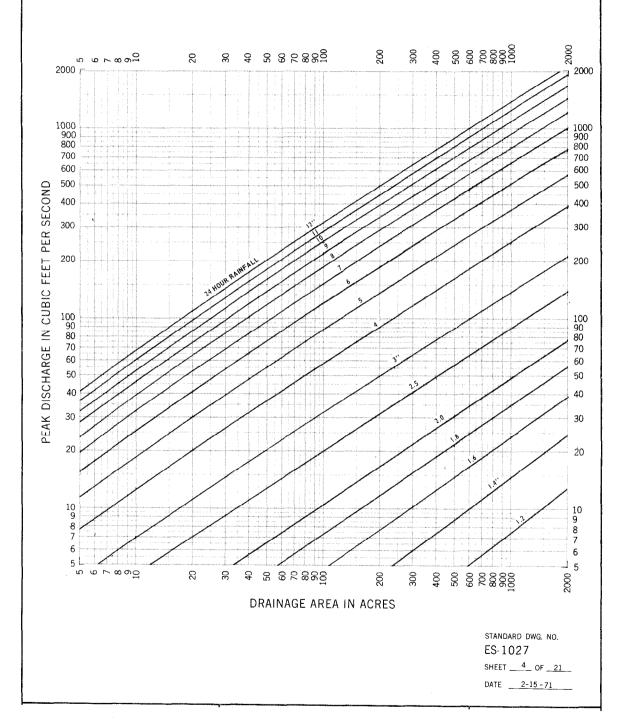
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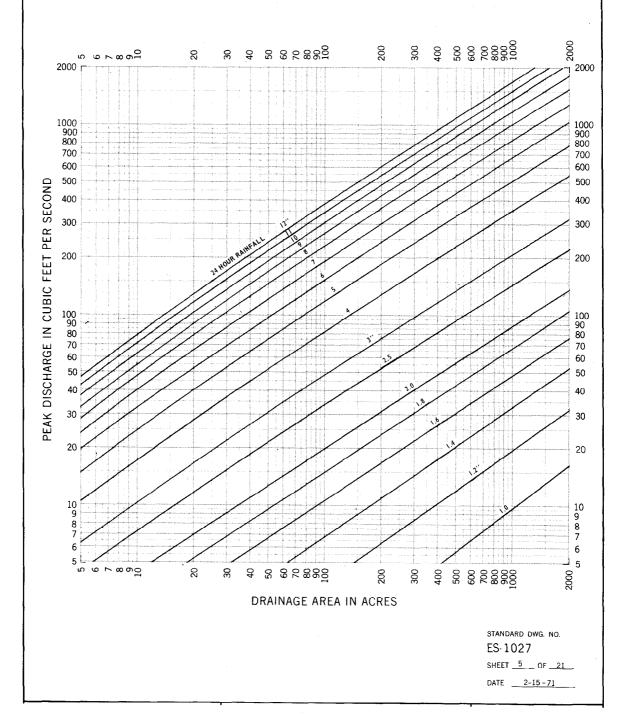
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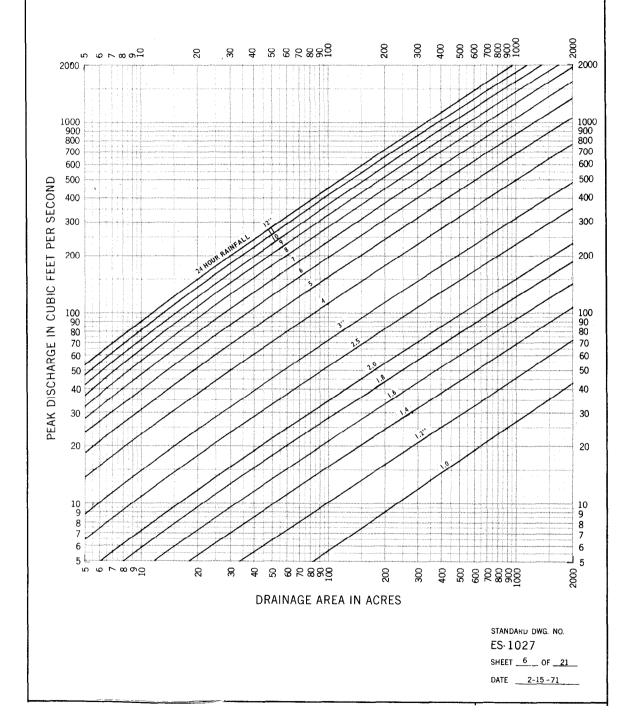
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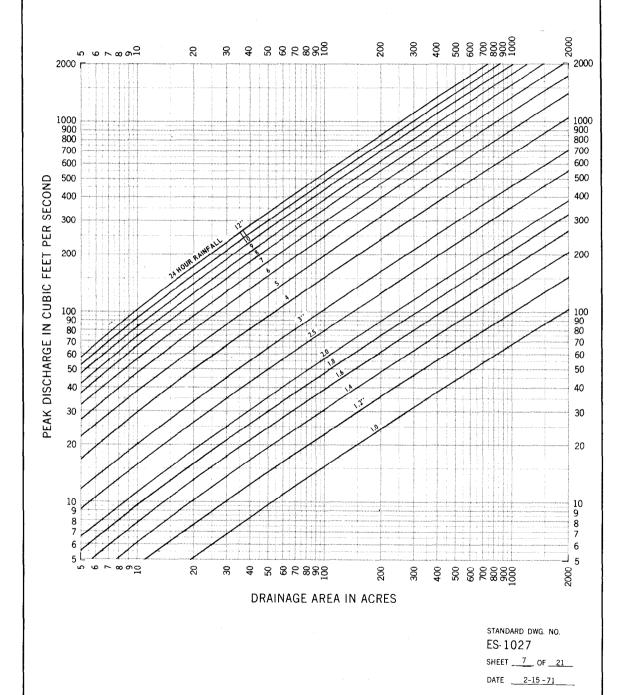
SLOPES - FLAT
CURVE NUMBER - 80



SLOPES - FLAT CURVE NUMBER - 85

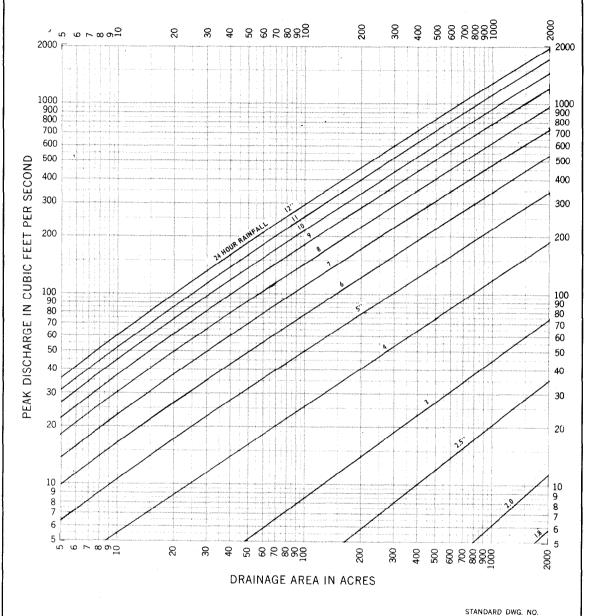


SLOPES - FLAT
CURVE NUMBER - 90



SLOPES - MODERATE
CURVE NUMBER - 60

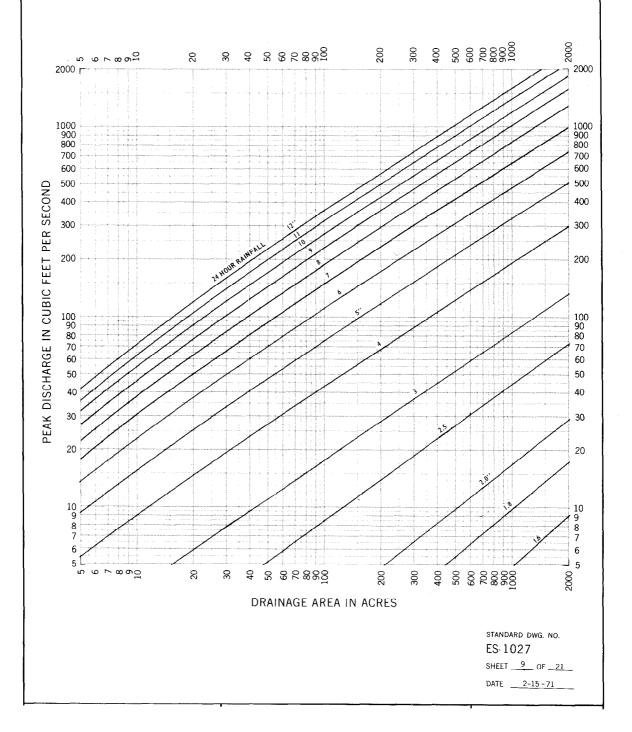
24 HOUR RAINFALL FROM US WB TP-40



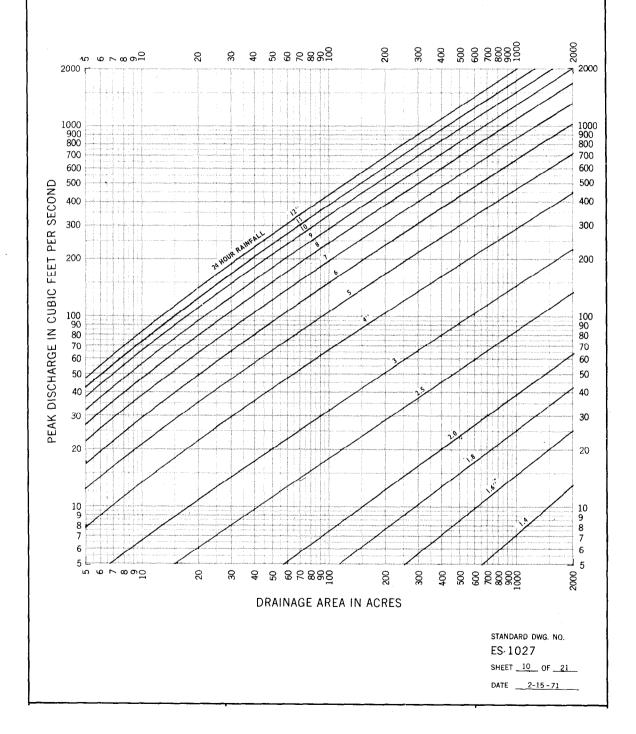
ES- 1027
SHEET 8 OF 21

DATE <u>2-15-71</u>

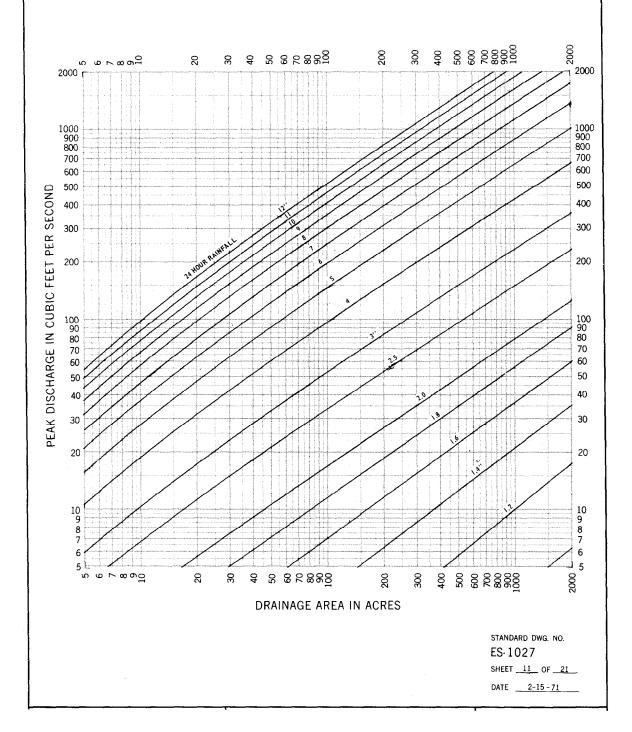
SLOPES - MODERATE
CURVE NUMBER - 65



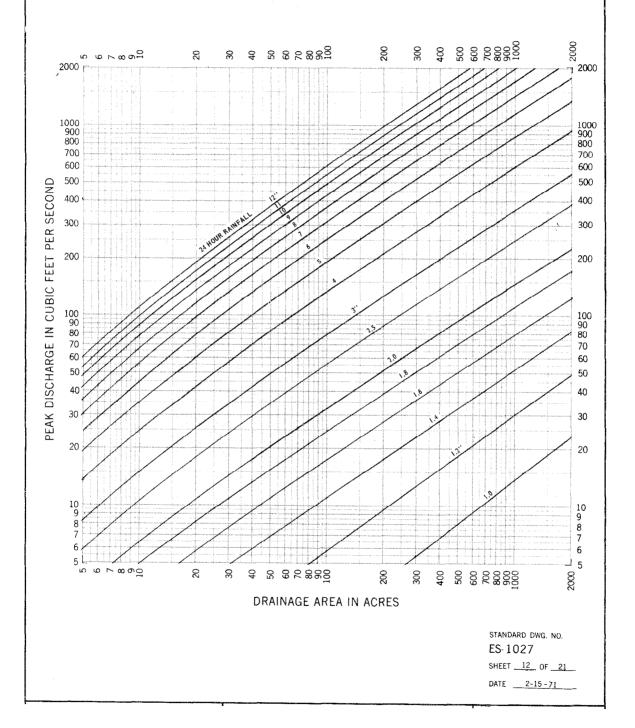
SLOPES - MODERATE CURVE NUMBER - 70



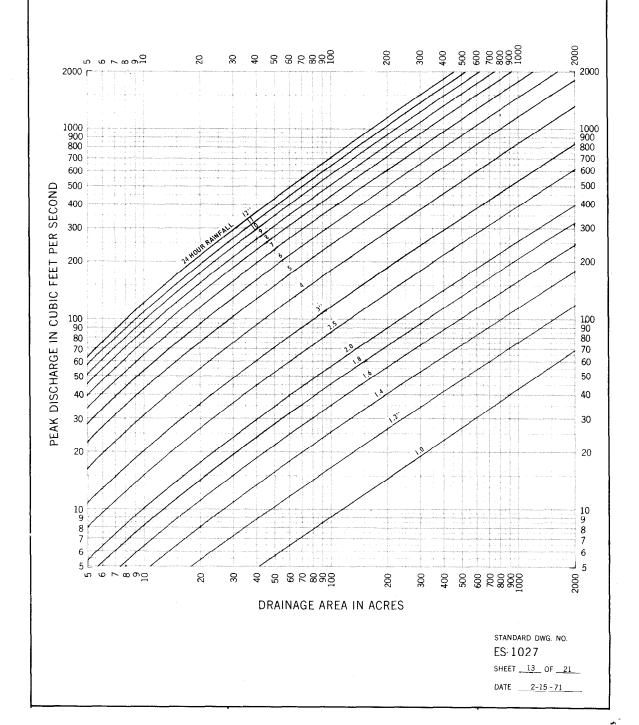
SLOPES - MODERATE
CURVE NUMBER - 75



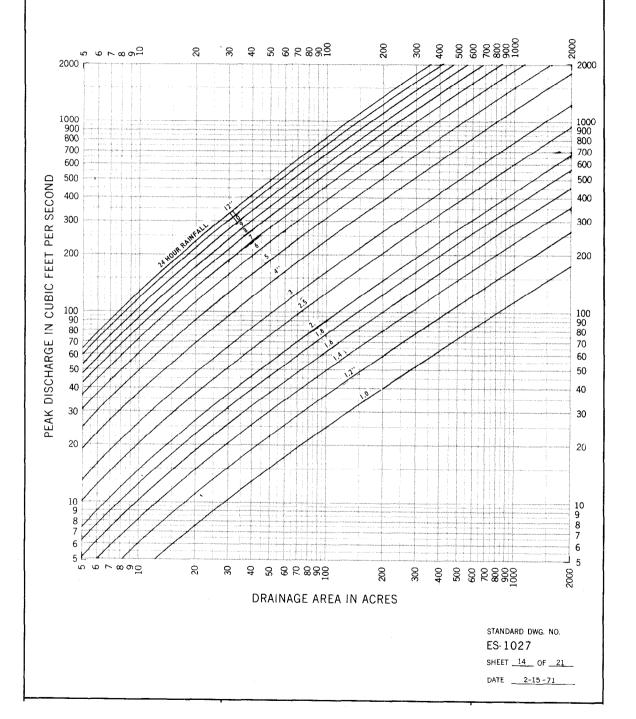
SLOPES - MODERATE CURVE NUMBER - 80



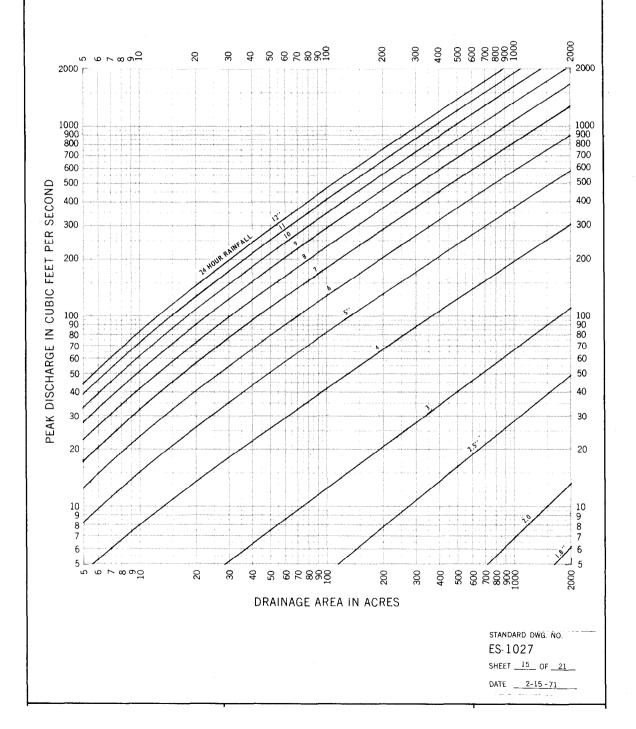
SLOPES - MODERATE
CURVE NUMBER - 85



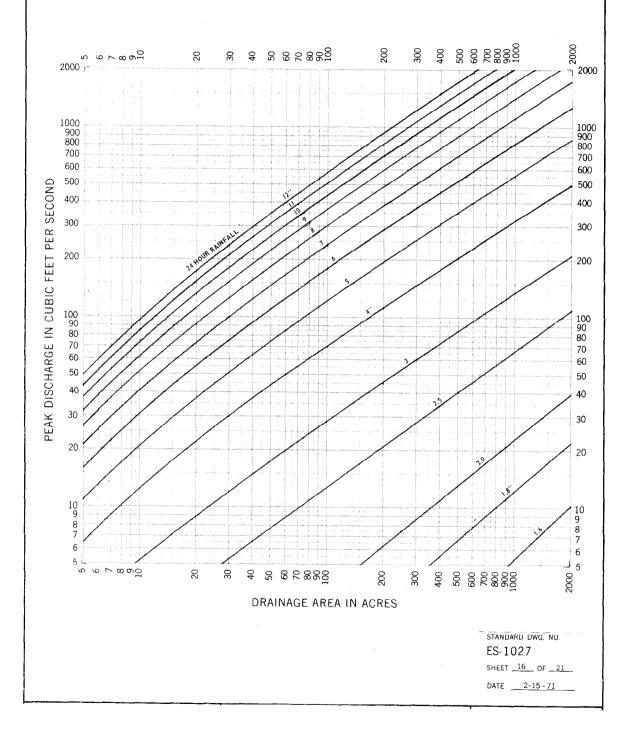
SLOPES - MODERATE CURVE NUMBER - 90



SLOPES - STEEP
CURVE NUMBER - 60



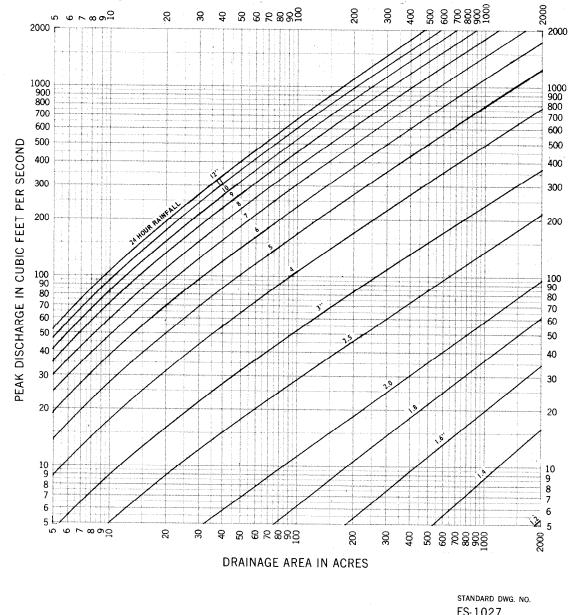
SLOPES - STEEP CURVE NUMBER - 65



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SLOPES - STEEP CURVE NUMBER - 70

24 HOUR RAINFALL FROM US WB TP-40



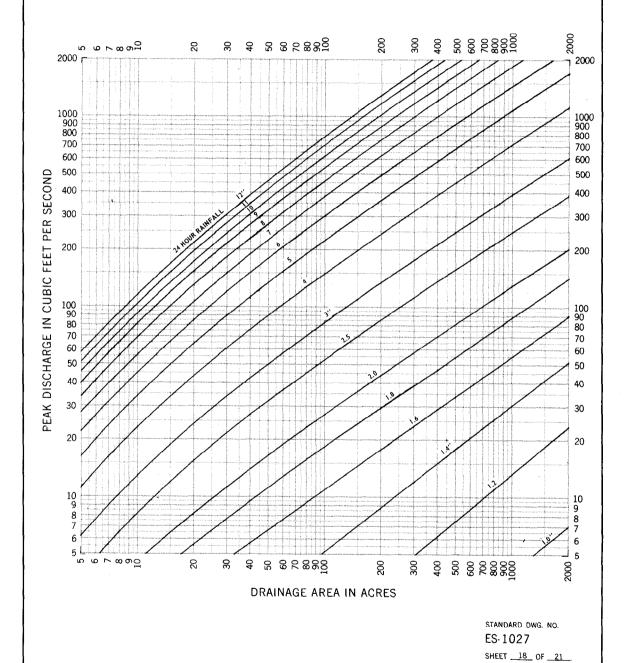
ES-1027

SHEET 17 OF 21

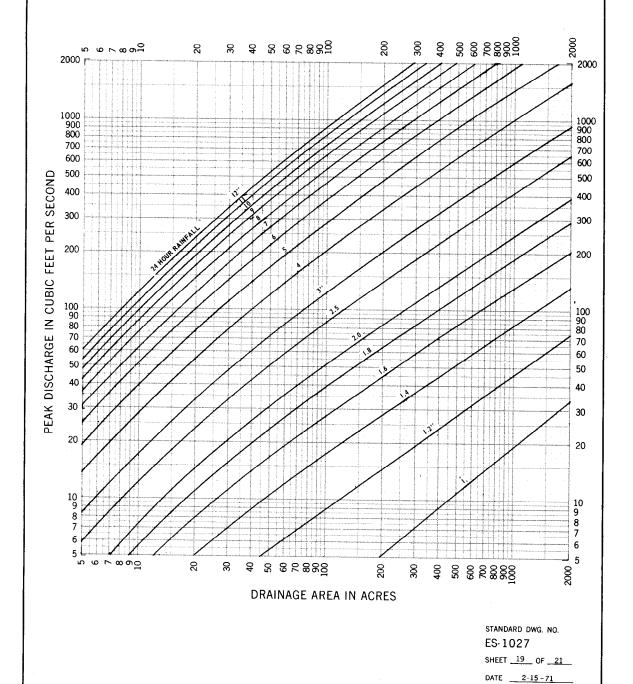
DATE 2-15-71

SLOPES - STEEP CURVE NUMBER - 75

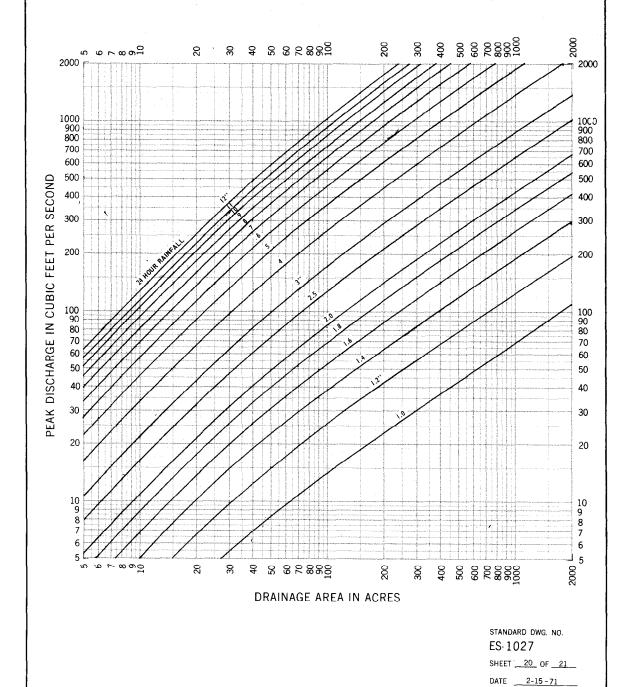
DATE 2-15-71



SLOPES - STEEP CURVE NUMBER - 80



SLOPES - STEEP
CURVE NUMBER - 85



SLOPES - STEEP CURVE NUMBER - 90

